

1 Supplementary Information for:

2 Plate motion drives variability in ocean oxygenation

3 through the Phanerozoic

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## 22 **Supplementary Discussion**

23 Most of the simulations in Extended Data Figs. 5 and 6 show that the ocean dynamics  
24 eventually settles to a stable steady-state solution. However, amongst these are cases where  
25 the system overshoots (for example, compare panels a and g of Extended Data Fig. 5) before  
26 approaching a steady-state solution. This implies the presence of an oscillatory mode, in  
27 agreement with the linear stability analysis of low-resolution GCM steady states under  
28 present-day conditions<sup>47,48</sup>. The simulations in Extended Data Figs. 5 and 6 reveal that,  
29 depending on changes to continental configuration, the steady-state solutions can lose  
30 stability, resulting in stable oscillations. It is known that, under certain climate conditions,  
31 advective feedbacks in the ocean circulation can destabilize the steady-state<sup>49</sup>. Extended Data  
32 Figs. 5bb and 6z are cases of steady-state solutions with oscillatory modes that are linearly  
33 stable, but only weakly stable, resulting in clear examples of damped oscillations. Since the  
34 frequencies of the damped and self-sustained oscillations of Extended Data Figs. 6z and 6aa,  
35 respectively, are very similar, the transition could represent a Hopf bifurcation, resulting in  
36 self-sustained oscillations around an unstable steady state.

37 The precise conditions that allow not only the destabilization of the steady state, but also  
38 the existence of *stable* oscillations, as well as the precise bifurcation structure associated with  
39 the oscillations, is the subject of ongoing work with idealized continental configurations.

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## 41 **Supplementary References**

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