

Comparison of Precipitation Projections of CMIP5 and CMIP6 Global Climate Models over Yulin, China

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6 **Comparison of Precipitation Projections of CMIP5 and CMIP6 Global Climate Models**
7 **over Yulin, China**

8
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23 **Comparison of Precipitation Projections of CMIP5 and CMIP6 Global Climate Models**
24 **over Yulin, China**

25
26 **Abstract**

27 This study compared precipitation projections of Coupled Model Intercomparison Project 5
28 (CMIP5) and 6 (CMIP6) GCMs over Yulin City, China. The performance of CMIP5 and CMIP6
29 GCMs in replicating Global Precipitation Climatology Centre (GPCC) precipitation climatology
30 of the city was evaluated using different statistical metrics. The best performing GCMs common
31 to both CMIP5 and CMIP6 were selected and subsequently downscaled to GPCC resolution
32 using linear scaling method to spatiotemporal changes in precipitation. The study revealed
33 BCC.CSM1.1(m), IPSL-CM5A-LR, MRI.CGCM3 and MIROC5 of CMIP5 and their equivalents
34 BCC-CSM2-MR, IPSL-CM6A-LR, MRI.ESM2.0 and MIRCO6 of CMIP6 as the most suitable
35 GCMs for the projection of rainfall in Yulin. Changes in precipitation were in the range of -14.0
36 – 0.0% and -22.0 – 0.2% during 2021–2060 for RCP4.5 and SSP2-4.5 respectively. The highest
37 decrease of -29.7 – -22.0% was projected by MRI-ESM-2-0 for SSP2-4.5, while -28.0 – -20.0%
38 by MIROC5 for RCP4.5. For RCP8.5 and SSP5-8.5, precipitation was projected to decrease in
39 the range of -17.0 – -2.0% and -32.0 – 0.0%, respectively during 2021 – 2060 by most of the
40 GCMs. An increase in precipitation up to 12.3% was projected only by IPSL-CM5A-LR for
41 RCP8.5 for this period. The highest decrease was projected by MIROC5 (-40.2 – -29.0%) for
42 RCP8.5 and IPSL-CM6A-LR (-40.2 – -26.0%) for SSP5-8.5. Overall, the results revealed a
43 higher decrease in precipitation in Yulin city by CMIP6 GCMs compared to those projected by
44 their corresponding GCMs of CMIP5 for both scenarios.

45 **Keywords:** CMIP6, climate change, global circulation model, statistical indices, Yulin

46

47 **1. Introduction**

48 Climate change has been an issue of concern over several decades due to its devastating
49 impacts. It has increased the risks of flooding (Manawi et al., 2020; Rahman et al., 2019),
50 occurrences of droughts (Alamgir et al., 2019; Ayugi et al., 2020; Shiru et al., 2019a), heatwaves
51 (Kang and Eltahir, 2018; Khan et al., 2019; Satyanarayana and Rao, 2020) and ecosystem
52 damages (Kim et al., 2019; Pérez-Ruiz et al., 2018). Many sectors including water resources,
53 agriculture, energy and health among others have also been affected by the changing climate.

54 Like many other parts of the globe, China is also experiencing the impacts of climate change.
55 Flooding is a common occurrence in different parts of China including cities like Yulin (Huang
56 et al., 2015; Yang et al., 2017) and it often causes damages to sectors such as power, agriculture,
57 health and sometimes leads to loss of lives (He et al., 2018). In 2003, heavy precipitation caused
58 river flooding leading to the destruction of 17,438 acres of farmlands, destruction of roads, and
59 damages to about 1,458 houses (He et al., 2018). Heavy precipitation affected more than 150,000
60 people in 2012 with 15 people missing and 11 deaths, destruction of roads and disruption of
61 aviation due to the heavy rain (China-Daily 2012). The disaster caused a direct economic loss in
62 the Yulin and Yan'an region up to 134 million yuan (\$21 million). Similarly, the flood in 2016
63 affected more than 21,000 people and evacuation of about 1200 persons (He et al., 2018). The
64 flood led to the damage of more than 850 houses, destruction of more than 1100 kilometers of
65 roads, damages to 30 bridges and culverts, and caused 11 landslides and other related geological
66 disasters. The economic implication of the disaster was about 150 million yuan. The region is
67 also known to be affected by droughts (Wang et al., 2020; Yin et al., 2020).

68 Understanding the expected changes in climate is crucial for the areas susceptible to
69 disasters in order to develop adaptation and mitigation plans against climate change. This is
70 particularly important using the recently released global climate models (GCM) of the Coupled
71 Model Inter-comparison Phase 6 (CMIP6). It is also crucial to assess how the CMIP6 differ in
72 projection from the previous Coupled Model Inter-comparison Phase 5 (CMIP5) in order to
73 streamline the existing adaptation measures based on CMIP5 projections (Jiang et al., 2020;
74 Song et al., 2021).

75 GCMs are developed by different institutions and have different performances in different
76 parts of the globe (Chen et al., 2017). Hence, the selection of the most realistic ones for a reliable
77 projection of climate for a region is crucial. The assessment of the performances of GCMs has

78 been conducted using different statistical indices (Rivera and Arnould, 2020; Sreelatha and
79 Anand Raj, 2019). However, due to contradictory outputs from different statistical measures,
80 supporting such outputs with other measures can be beneficial in reaching a compromise.

81 As GCMs are characterized by coarser spatial resolutions, their applications in climate
82 projections and climate change impact studies at local and regional scales can be unreliable
83 (Onyutha et al., 2016; Pour et al., 2018; Salman et al., 2018). Therefore, they are required to be
84 downscaled (Almazroui et al., 2020a; Sa'adi et al., 2019; Shiru et al., 2020) using either the
85 dynamical downscaling or the statistical downscaling (SD) methods. The SD method is known to
86 have the advantages of computational efficiency and cost-effectiveness, the possibility of
87 incorporation of observations directly into methods, and provision of point-scale climatic
88 projections from GCM-scale (Fowler et al., 2007).

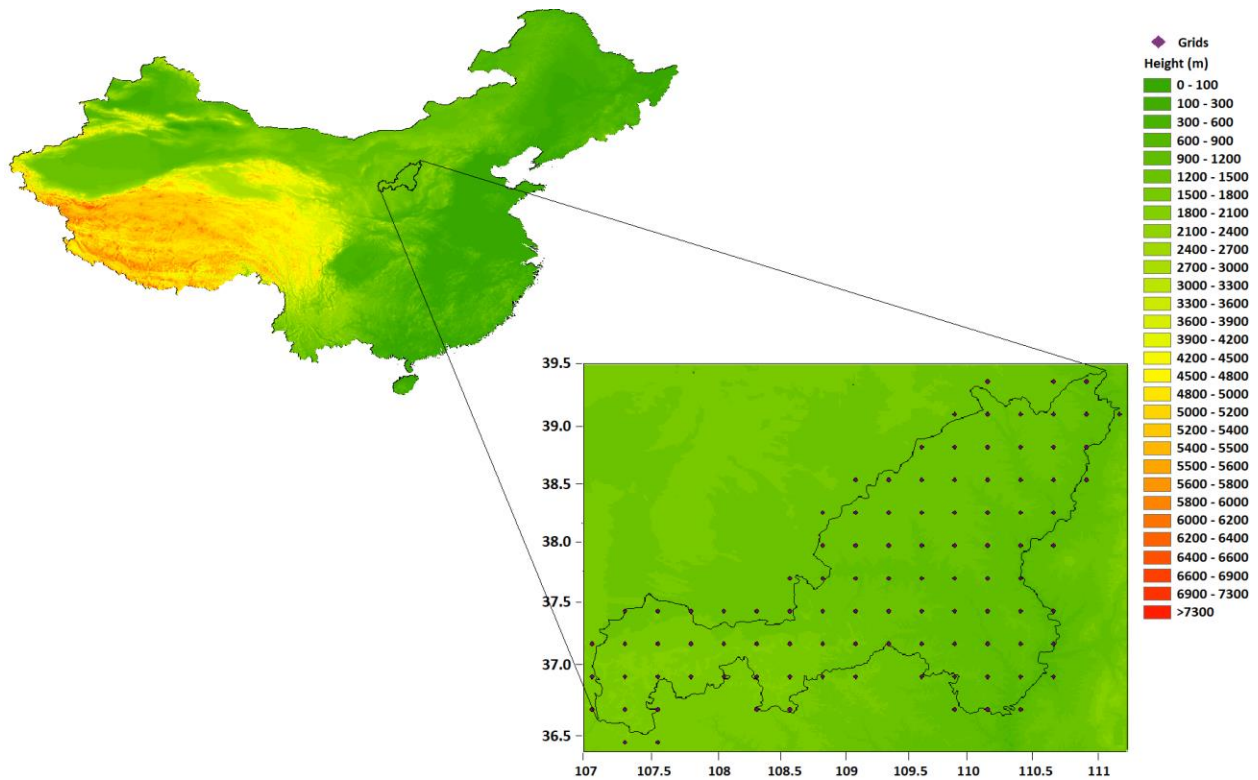
89 An array of statistical metrics was used to select the best performing GCMs for the
90 projection of precipitation in the study area. Equivalent GCMs of the CMIP5 and CMIP6 for
91 selected for the comparison of projections. The selected GCMs were downscaled using linear
92 scaling method. This study employs only the GCMs which can reliably simulate the exiting
93 climate of the study area and thus, capable to provide a trustworthy comparison of CMIP5 and
94 CMIP6 projects. It is expected that the comparison of precipitation projections of CMIP5 and
95 CMIP6 would help in streamlining the existing adaptation measures formulated based on CMIP5
96 projections or deriving new measures based on new scenarios of CMIP6.

97

98 **2. Study Area and Datasets**

99 *2.1. Study Area*

100 The study area, Yulin (Figure 1) is located in the northern Shaanxi province of China (Longitude:
101 107°15'47"–111°14'44"E; Lattitude: 36°49'07"–39°34'47"N). Yulin covers a total area of 385 km
102 by 263 km (Zha, et al. 2008). The terrain of the area descends from 1,907 m in the east to 585 m
103 in the west above the mean sea level. Yulin has a semi-arid temperate continental monsoon type
104 climate which is characterized by dry and little precipitation in spring and winter and high
105 precipitation during summer and autumn. The annual average precipitation in Yulin is around
106 450 mm. The annual average temperature in the area is 9.6°C (Wang, 2016).



107

108

Figure 1 Map of Yulin showing the grid points considered for projection in the study.

109

2.2. Gauged based gridded precipitation

110

The GPCC full data reanalysis product of the Deutscher Wetterdienst (Becker et al., 2013;

111

Schneider et al., 2014) was used in this study as the reference data. The GPCC precipitation

112

amongst most other precipitation products has the advantages of (1) good data quality for

113

hydrological studies, (2) availability for a longer period, (3) better performance as being

114

developed using the highest number of collected precipitation records, and (4) completeness of

115

time series for the recent decades (Ahmed, et al. 2017; Spinoni, et al. 2014). This study used the

116

monthly precipitation data for the period 1961 – 2005. Data were collected from a total of 100

117

grid points to cover the whole Yulin. The location of the grid points is shown in Figure 1.

118

119

2.3 Global Climate Models

120

This study uses the historical and future simulations of GCMs of CMIP5 and CMIP6. The

121

CMIPs are sets of globally coordinated GCM simulations which comprises of historical and

122

future climate simulations assembled from different climate modeling groups. The CMIP5 offers

123 significant improvements compared to the CMIP3 (Taylor et al., 2012). The CMIP5 comprises of
 124 four scenarios called representative concentration pathways (RCPs). In the CMIP6, the four RCP
 125 scenarios of CMIP5, RCP2.6, RCP4.5, RCP6.0 and RCP8.5 have been updated as Shared
 126 Socioeconomic Pathways (SSPs) scenarios, SSP1-2.6, SSP2-4.5, SSP4-6.0, and SSP5-8.5
 127 respectively. Each GCM also considers 2100 radiative forcing levels. The GCMs of CMIP6 are
 128 developed through improved emission scenarios, land use data, physical processes and model
 129 parameterization to provide more realistic projections of future climate (Eyring, et al. 2016;
 130 O'Neill, et al. 2016). In this study, historical and future simulation of 10 GCMs of CMIP5 and
 131 CMIP6 were considered. The GCMs were chosen based on their availability from the same
 132 institution. The first ensemble members for both CMIP5 and CMIP6 were considered. The
 133 GCMs chosen for this study are provided in Table 1.

134

135 Table 1. Descriptions of the GCMs of CMIP5 and CMIP6 used in this study.

Institution - Country	CMIP5 Model Name	Resolution (lon/lat in °)	CMIP6 Model Name	Resolution (lon/lat in °)
Australian Community Climate and Earth System Simulator - Australia	ACCESS1.3	1.9 × 1.2	ACCESS.ESM1.5	1.9×1.2
Beijing Climate Center - China	BCC.CSM1.1(m)	1.125 × 1.125	BCC.CSM2.MR	1.1× 1.1
Canadian Centre for Climate Modelling and Analysis - Canada	CanESM2	2.8 × 2.8	CanESM5	2.8×2.8
NASA Goddard Institute for Space Studies - United States	GISS.E2.R	2.5 × 2.0	GISS.E2.1.G	2.5 × 2.0
Marchuk Institute of Numerical Mathematics - Russia	INM.CM4	2 × 1.5	INM.CM4.8	2×1.5
Institut Pierre-Simon Laplace - France	IPSL.CM5A.LR	3.75 × 1.875	IPSL.CM6A.LR	2.5× 1.3
The University of Tokyo, National Institute for Environmental Studies, and Japan Agency for Marine-Earth Science and Technology - Japan	MIROC5	1.4 × 1.4	MIROC6	1.4× 1.4
Max-Planck-Institut für Meteorologie - Germany	MPI.ESM.LR	1.9 × 1.9	MPI.ESM1.2.LR	1.9× 1.9
Meteorological Research Institute - Japan	MRI.CGCM3	1.25 × 1.25	MRI.ESM2.0	1.1× 1.1
Norwegian Meteorological Institute - Norway	NorESM1.M	2.5 × 1.875	NORES2.MM	2.9 x 1.9

136

137 3. Methodology

138 The methods used in this study are described in this section. For an unbiased comparison of
 139 model performance, the GPCC and the precipitation simulations of all GCMs were re-gridded to

140 a uniform resolution of $1^\circ \times 1^\circ$ (latitude \times longitude) using bilinear interpolation to have a uniform
 141 resolution. Bilinear interpolation is often conducted for transforming spatially coarse GCM data
 142 into finer data through GCM data interpolation from the four nearest neighboring grid points
 143 (Ahmed, et al. 2019; Almazroui, et al. 2020a). After selection of GCMs, the selected GCMs and
 144 the GPCC precipitation data were re-gridded to 0.25° and used for spatiotemporal projection of
 145 precipitation.

146 3.1 Statistical Indices

147 The ability of the different GCMs in reproducing the properties of GPCC precipitation at all the
 148 100 grid points of the study area was assessed using four statistical indices: normalized root
 149 mean square error (NRMSE), Percentage of Bias (Pbias), Volumetric Efficiency (VE), and
 150 Coefficient of Determination (R^2). Besides, probability density function (pdf) and spatial
 151 relationship of the mean monthly precipitation of the different GCMs were compared with
 152 GPCC precipitation to assess their abilities in replicating the precipitation climatology of the
 153 study area. Details about the statistical metrics used in this study are as follows. The expressions
 154 used to describe statistical metrics used here: $x_{pred,i}$ and $x_{obs,i}$ are the i -th GCM and GPCC data.

155 The magnitude of the errors in estimation can be summarized by NRMSE (Willmott 1982). It is
 156 a normalized statistic that provides a relative magnitude of the residual variance to the variance
 157 of the measured data. Smaller NRMSE values (preferably zero) indicate better performance of
 158 the model. NRMSE is defined as follows

$$159 \quad NRMSE = \frac{\left[\frac{1}{n} \sum_{i=1}^n (x_{pred,i} - x_{obs,i})^2 \right]^{1/2}}{\frac{1}{n} \sum_{i=1}^n x_{pred,i}} \quad (1)$$

160 The tendency of GCM to underestimate or overestimate the GPCC precipitation is measured
 161 using Pbias. Model performance is better when the Pbias is closer to zero. The PBIAS is a
 162 statistical metric that gives the estimate of over estimation or under prediction of a model
 163 (Wagena, et al. 2018). The evaluation of Pbias is conducted as follows.

$$164 \quad Pbias = 100 \times \left[\frac{\sum_{i=1}^n (x_{pred,i} - x_{obs,i})}{\sum_{i=1}^n x_{pred,i}} \right] \quad (2)$$

165 The VE measures the ratio between GCM and GPCC precipitation volumes over a period, where
 166 a VE value of 1 indicates a perfect estimation. It can be calculated using the following equation.

167
$$VE = 1 - \frac{\sum_{i=1}^n (x_{pred,i} - x_{obs,i})}{\sum_{i=1}^n x_{obs,i}} \quad (3)$$

168 The R^2 can be defined as the square of the Pearson's product-moment correlation coefficient (i.e.
 169 $R^2 = r^2$) describing the proportion of the total variance in the GPCC precipitation which is
 170 explainable by GCM (Legates and McCabe Jr 1999). R^2 values can range between -1.0 and 1.0,
 171 in which the higher absolute value indicates a better agreement. Computation of R^2 is as follows.

172
$$R^2 = \frac{\sum_{i=1}^N (x_{obs,i} - \overline{x_{obs}})(x_{Pred,i} - \overline{x_{Pred}})}{\sqrt{\sum_{i=1}^N (x_{Pred,i} - \overline{x_{Pred}})^2 \sum_{i=1}^N (x_{obs,i} - \overline{x_{obs}})^2}} \quad (4)$$

173
 174 *3.2 Downscaling of precipitation of selected GCMs of CMIP5 and CMIP6*
 175 The linear scaling (LS) method was applied for the downscaling of the selected GCMs. LS
 176 (Lenderink, et al. (2007) uses the monthly correction values obtained from the difference in
 177 GPCC and GCMs simulated monthly precipitation for the reference period. The monthly scaling
 178 factor is then applied to raw GCM data. The monthly precipitation, P is corrected using the
 179 following equation:

180
$$P^* = \alpha P \quad (5)$$

181 where $\alpha = P_o/P_s$ (6)

182
 183 P_o is the monthly mean of GPCC precipitation whereas P_s is the monthly mean of the GCM
 184 simulated precipitation. LS method requires less information such as only monthly data to
 185 calculate the scaling factor (Lafon, et al. 2013) and thus, widely used for precipitation
 186 downscaling.

187
 188
 189
 190 **4. Results**

191 *4.1 Performance assessment of GCMs using statistical metrics*

192 The results of the statistics used for the performance evaluation of different historical GCMs are
 193 presented in Table 2. It shows a variation in the performances of different GCMs. For CMIP5,

194 the GCMs with the best performance metrics are ACCESS1.3, BCC.CSM1.1(m), IPSL.CM5A.LR,
 195 and MIROC5. The GCMs of the same institutions also showed good performances for CMIP6
 196 except for ACCESS.ESM1.5 which showed a relatively poor performance.

197

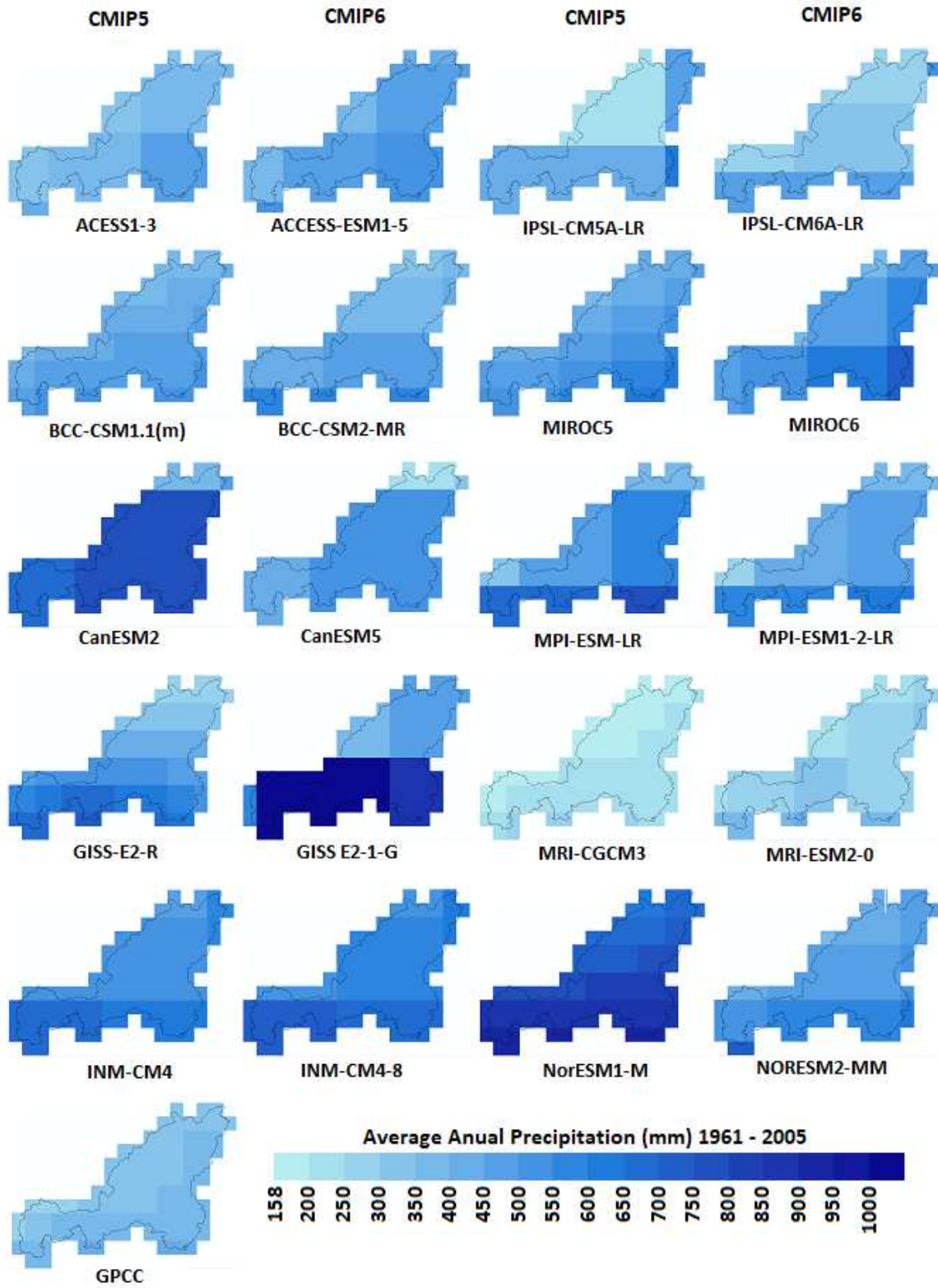
198 Table 2. Statistical metrics showing the performances of GCMs of CMIP5 and CMIP6 in replicating GPCC
 199 precipitation over Yulin (Best metrics are presented in bold).

CMIP5					CMIP6				
GCM	NRMSE	Pbias	VE	R ²	GCM	NRMSE	Pbias	VE	R ²
ACCESS1.3	76.3	10.9	0.68	0.73	ACCESS.ESM1.5	154.4	23.8	0.67	0.77
BCC.CSM1.1(m)	144.8	21.2	0.79	0.79	BCC.CSM2.MR	115	19.8	0.8	0.63
CanESM2	292.9	49.8	0.5	0.16	CanESM5	136.1	23.7	0.73	0.46
GISS.E2.R	135.8	29.1	0.68	0.22	GISS.E2.1.G	175.8	53.3	0.47	0.04
INM.CM4	274.7	38.8	0.61	0.46	INM.CM4.8	279.9	41.5	0.59	0.45
IPSL.CM5A.LR	81.0	-4.5	0.77	0.36	IPSL.CM6A.LR	70.2	6.1	0.8	0.59
MIROC5	177.2	27.5	0.72	0.76	MIROC6	181.6	23.3	0.72	0.65
MPI.ESM.LR	174.3	35.5	0.65	0.41	MPI.ESM1.2.LR	123.3	23.6	0.76	0.54
MRI.CGCM3	173.8	-44.6	0.70	0.85	MRI.ESM2.0	85.8	-21.3	0.78	0.78
NorESM1.M	397.3	55.1	0.45	0.21	NORES2.MM	182.6	29.8	0.7	0.62

200

201 *4.2 Spatial relationship between GCMs and GPCC precipitations*

202 The ability of different GCMs in replicating the spatial distribution of GPCC precipitation for the
 203 study area are presented in Figure 2. The performances of the GCMs were found to vary widely
 204 in reproducing the GPCC precipitation. Among CMIP5 GCMs, the highest overestimations were
 205 by CanESM2 and NorESM1-M, while an underestimation was by MRI-CGCM3 followed by
 206 IPSL-CM5A-LR in some parts. For CMIP6, GISS-E2-1G showed the highest overestimation of
 207 precipitation. Generally, the GCMs with better performance metrics (Table 2) showed better
 208 skills in replicating the precipitation climatology of GPCC for the study area.



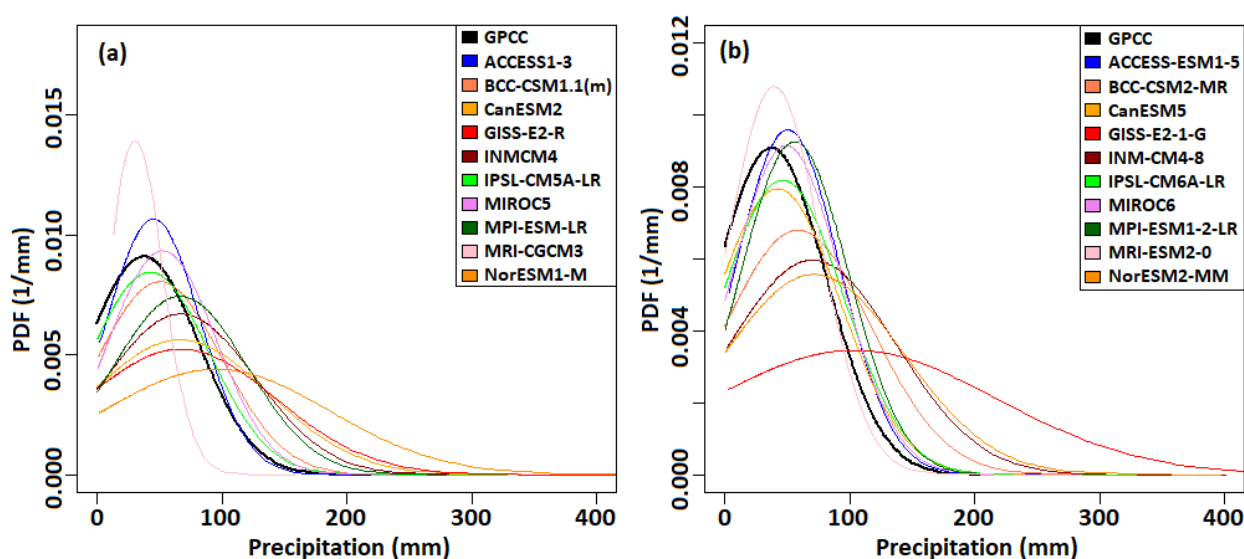
209

210 Figure 2 Spatial distribution of average annual precipitation of GCMs of CMIP5 and CMIP6 compared to
 211 that of GPCC during 1961–2005.

212

213 4.3 Comparison using probability density function

214 The PDFs of monthly GCM precipitation was compared with the PDF of GPCC precipitation for
 215 the study area. The results for CMIP5 and CMIP6 are presented in Figure 3(a) and 3(b)
 216 respectively. Figures show that most GCMs were able to replicate the precipitation properties of
 217 the GPCC, especially the skewness. However, the distribution of precipitation was found better
 218 for the GCMs which showed a better performance in terms of statistical metrics presented in
 219 Table 2.



220

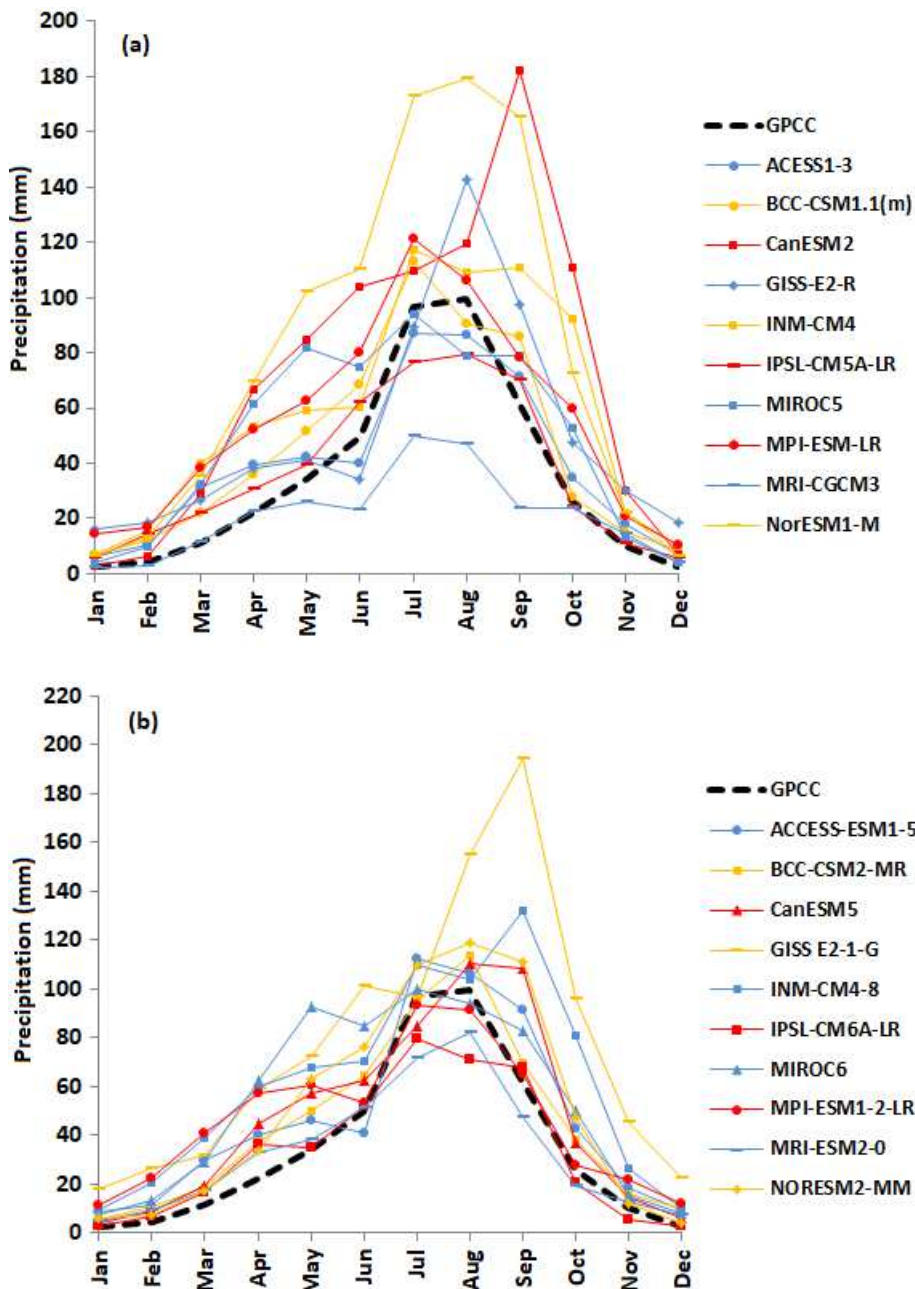
221 Figure 3 Probability density function of monthly precipitations of GPCC and the GCMs of (a) CMIP5 and
 222 (b) CMIP6.

223

224 4.4 Comparison of GCMs in reproducing monthly mean precipitation

225 The mean monthly GCM precipitation was compared with mean monthly GPCC
 226 precipitation for the period 1961–2005. Obtained results for the GCMs of CMIP5 and CMIP6
 227 are presented in Figure 4(a) and (b) respectively. There was a variation in the estimation of
 228 GPCC precipitation by the GCMs, especially during the wet season. CanESM2 and NorESM1-M
 229 of CMIP5 overestimated the precipitation for all the months, while an underestimation was by
 230 MRI-CGCM3. For the CMIP6, overestimation was by GISS-E2-1-2G and underestimation by
 231 MRI-ESM2-0 and IPSL-CM5A-LR, especially during the wet period. Though most of the

232 CMIP6 GCMs were found to overestimate the GPCC precipitation, overall they were found more
 233 capable in replicating the mean monthly precipitation of GPCC compared to CMIP5 GCMs.



234

235

236 Figure 4 Comparison of mean monthly precipitation of GCMs of (a) CMIP5 and (b) CMIP6 with that of GPCC
 237 precipitation.

238

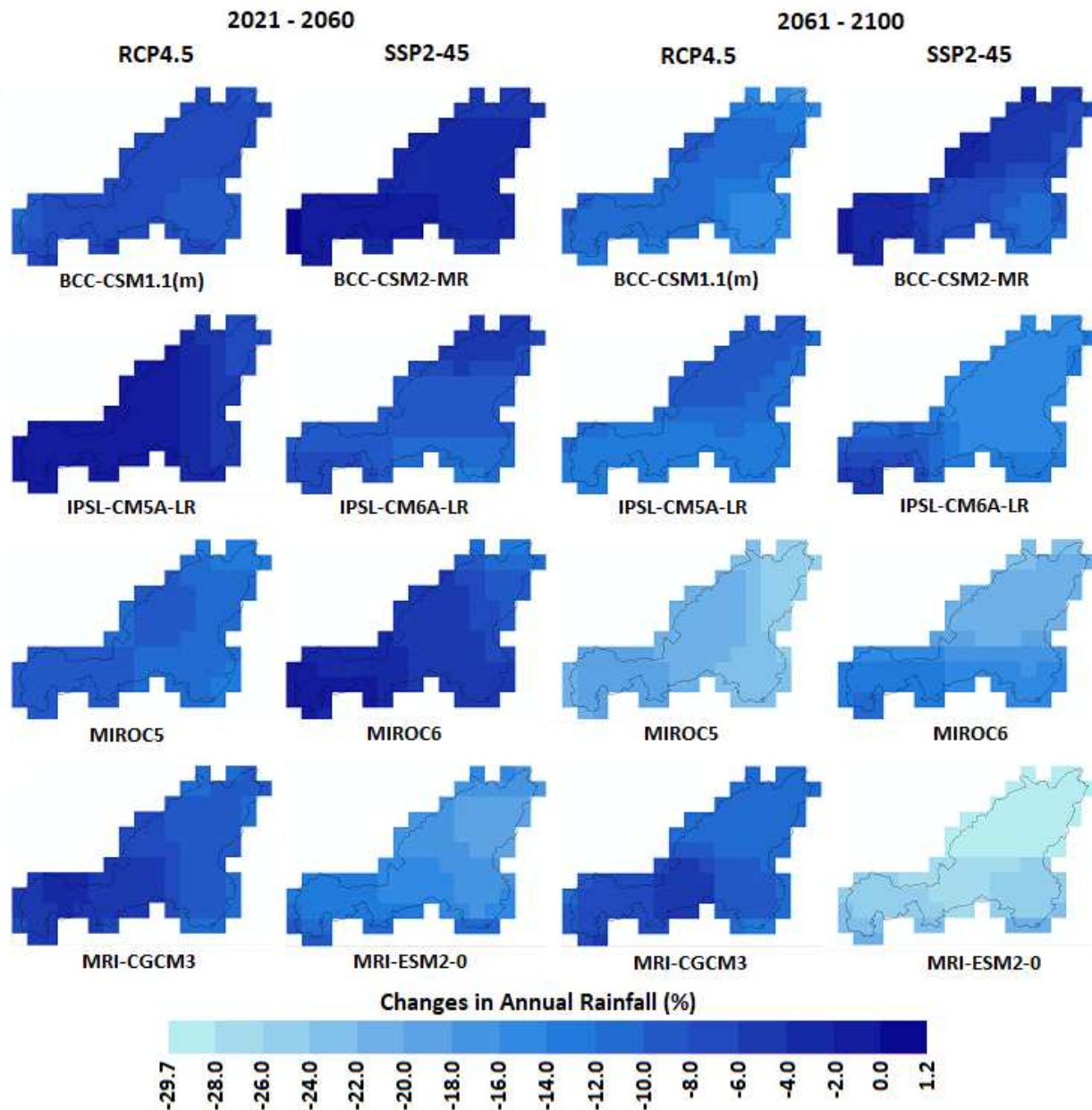
239 *4.5 Selection of the best performing GCMs*

240 The performance of the GCMs based on statistical indices and replication of PDF, and spatial
241 precipitation distribution patterns were considered. Based on the statistics, BCC.CSM1.1(m),
242 IPSL.CM5A.LR, and MIROC5 were the better performing GCMs for both CMIP5 and CMIP6.
243 Besides, ACCESS1.3 showed overall a better performance compared to the rest of the CMIP5
244 GCMs, while its equivalent under CMIP6 showed a poor performance than MRI.ESM2.0. As
245 CMIP6 is a more recent simulation, the MRI.CGCM3 for CMIP5 and its equivalent GCM for
246 MICP6, MRI.ESM2.0 were prioritized as the fourth model for projection of precipitation.

247

248 *4.7 Projection of precipitation from the selected GCMs of CMIP5 and CMIP6*

249 The spatial projections of precipitation for the study area by CMIP5 GCMs for RCP 4.5 and
250 CMIP6 GCMs for SSP2-4.5 for two future periods, 2021–2060 and 2061 – 2100 are presented in
251 Figure 5. Large heterogeneity in precipitation changes was projected by different GCMs for
252 RCPs and SSPs and the two projection periods. During 2021 – 2060, the highest decrease in
253 precipitation was projected by MRI-ESM-2-0 for SSP2-4.5 while the highest decrease for RCP
254 4.5 was projected by MIROC5. For the same period, BCC-CSM2-MR projected an increase in
255 precipitation by 1.2% at the extreme west of the study area. Percentage change in precipitation
256 was in the range of -14.0 – 0.0% for RCP4.5 while it was -22.0 – 1.2% for SSP2-4.5. During
257 2061 – 2100, all the GCMs projected decreases in precipitation, with the highest decrease of -
258 29.7 – -22.0% by MRI-ESM-2-0 for SSP2-4.5 and -28.0 – -20.0% by MIROC5 for RCP4.5. The
259 lowest decrease during this period was projected by MRI-CGCM3 and BCC-CSM2-MR for
260 RCP4.5 and SSP2-4.5 respectively.

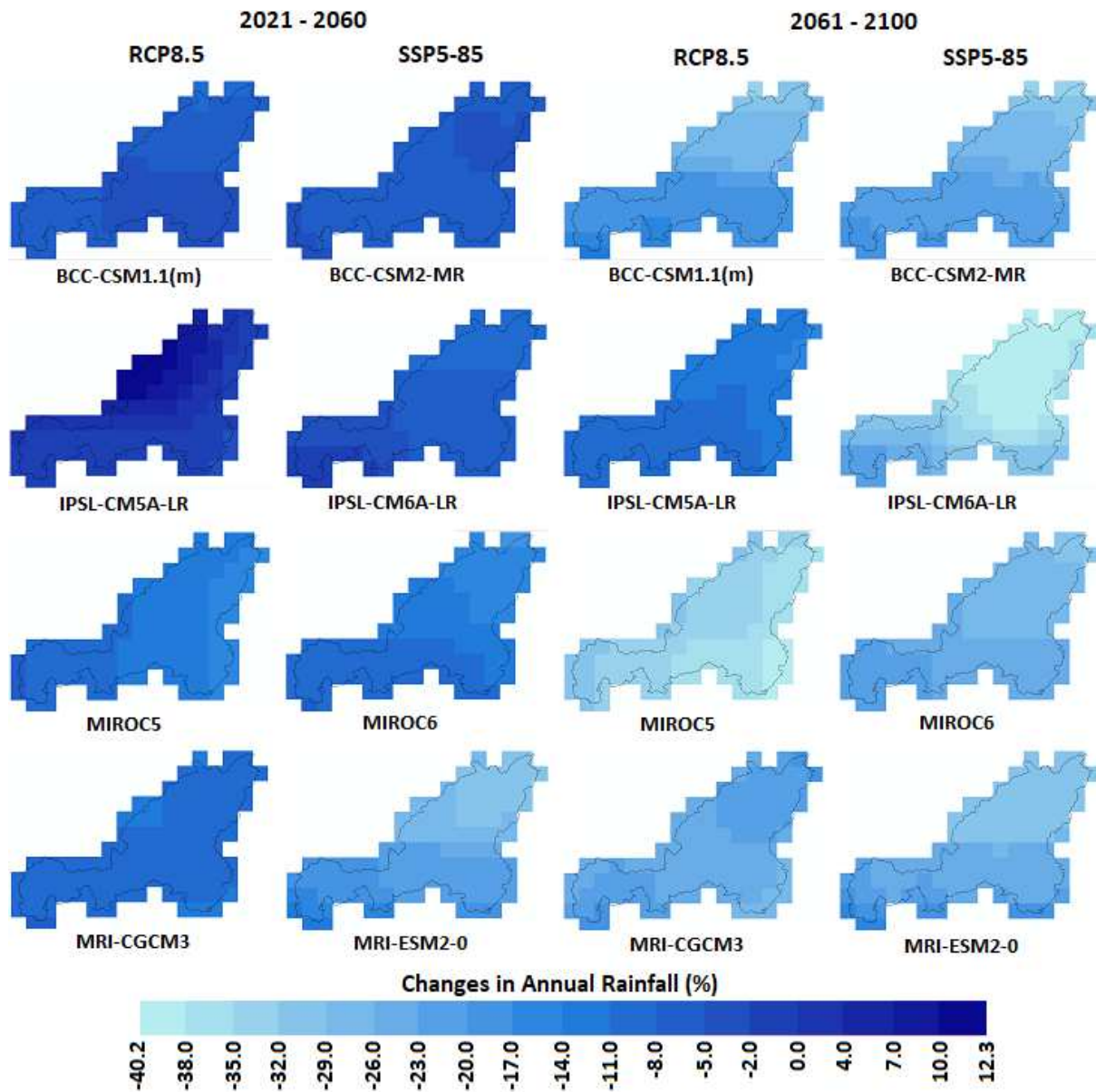


261

262 Figure 5. Spatial distribution of projected precipitation by the GCMs of CMIP5 and CMIP6 during 2021–
 263 2060 and 2061–2100 for RCP4.5 and SSP2-4.5.

264 The spatial distribution of projected precipitation in the study area by CMIP5 GCMs for
 265 RCP 8.5 and CMIP6 GCMs for SSP5-8.5 for the periods 2021–2060 and 2061–2100 are
 266 presented in Figure 6. Compared to RCP 4.5 and SSP2-4.5, RCP8.5 and SSP5-8.5 showed higher
 267 decreases in precipitation. During 2021 – 2060, the projected decrease in precipitation was in the
 268 range of -17.0 – -2.0% by the CMIP5 GCMs for RCP8.5 while the decrease was projected

269 between -32.0 and 0.0% by the CMIP6 GCMs for SSP5-8.5. Increases in precipitation of 0.0 –
 270 12.3% were noticed only for IPSL-CM5A-LR for RCP8.5. During 2061 – 2100, the decrease in
 271 precipitation was projected the highest by MIROC5 (-40.2 – -29.0%) for RCP8.5 and IPSL-
 272 CM6A-LR (-40.2 – -26.0%) for SSP5-8.5.



274 Figure 6. Spatial distribution of projected precipitation by the GCMs of CMIP5 and CMIP6 during 2021–
 275 2060 and 2061–2100 for RCP8.5 and SSP5-8.5.

276

277 **5. Discussion**

278 Climate change remains a major challenge in many parts of the globe as they have
279 devastating impacts on several sectors. Many studies have shown that the expected changes in
280 climate will result in increased temperatures and more erratic precipitations in many parts of the
281 world. Projection of precipitation under CMIP5 in Nigeria showed that while precipitation will
282 increase in some parts of the country, particularly the semi-arid and arid regions where they are
283 usually low, the other parts where they used to be higher will experience decreases (Shiru, et al.
284 2019b). (Homsy, et al. 2020) used CMIP5 GCMs in Syria and showed that precipitation would
285 increase by up to 87% in some parts and decrease by up to -85% in the coastal areas. Many other
286 studies have also shown both increase and decrease in precipitation in many parts of the world
287 using CMIP5 GCMs (Iqbal, et al. 2020; Narsey, et al. 2020; Shiru and Park 2020; Ullah, et al.
288 2020).

289 The studies for the recently released CMIP6 GCMs have also shown such changes in
290 precipitation. The study conducted over South Asian countries Almazroui, et al. (2020c) showed
291 that the annual mean precipitation will increase by 27.3% in India, 18.9% in Bhutan, 26.4% in
292 Pakistan, 19.5% in Nepal, 25.1% in Sri Lanka and 17.1% in Bangladesh in the last part of the
293 century under SSP5-8.5 scenario. Over Africa, (Almazroui, et al. 2020b) projected precipitation
294 under CMIP6 and showed that while the northern and the southern parts of Africa are analyzed
295 to witness a reduction in precipitation, the central parts are expected to have increases of 6.2%,
296 6.8%, and 9.5% for SSP1-2.6, SSP2-4.5, and SSP5-8.5, respectively.

297 In China, studies showed variations in the projected precipitations for both CMIP5 and
298 CMIP6 with some studies showing mostly increases while others showed a simultaneous
299 increase and decrease depending on the region. The annual precipitation was projected to
300 increase significantly relative to the present day for CMIP5 (Chen 2013). This study revealed
301 that the increase of precipitation in the Northwest of China is primarily due to the increase in
302 light showers while the increases in precipitation in the north and northeastern parts are due to an
303 increase in medium precipitation. The increases in precipitation are expected in the southern
304 parts of China due to an increment in heavy precipitation. Zhou, et al. (2019) showed both
305 increases and decreases in daily precipitation under different warming conditions over China
306 using CMIP5. The increase in warming of 1.5°C was projected to cause an increase in the

307 frequency and intensity of precipitation in northeast China, north China, and the Qinghai–Tibet
308 Plateau while there would be a decrease in total precipitation in south China and southwest
309 China. The number of wet days were projected to increase in the north while the decreases in the
310 south under 2°C.

311 The projection of the changes in precipitation over northwestern China for CMIP6 (Su-
312 Yuan, et al. 2020) showed that there would be lesser warming compared to that previously
313 expected, which would affect the patterns in precipitation changes. Unlike in the historical period
314 (1850 – 2014) when the rate of warming was 0.05°C per decade, the annual mean temperature is
315 projected to increase up to 0.06°C, 0.26°C and 0.59°C per decade for SSP1-2.6, SSP2-4.5 and
316 SSP5-8.5, respectively for the period 2015–2099. The total annual precipitation for the area is
317 projected to increase by 5.6, 6.4 and 8.0 mm/decade for SSP1-2.6, SSP2-4.5 and SSP5-8.5,
318 respectively.

319 In this study, both increases and decreases in precipitation are projected for the study
320 area. Zhou, et al. (2019) projected the increases in precipitation over China, except some
321 northwestern parts where Yulin belongs, which supports the findings of this study. The projected
322 the increases in some parts of the study area are also corroborated by other studies.

323

324 **6. Conclusions**

325 This study compares the projections of precipitation by CMIP5 and CMIP6 GCMs over
326 Yulin city of China. Different statistical and graphical metrics were used in assessing the ability
327 of 10 GCMs in replicating the precipitation properties of the study area. Finally, four GCMs
328 having the highest ability in replicating the properties of GPCP precipitation were selected for
329 the projection of precipitation over the study area. This study revealed that ACCESS1.3,
330 BCC.CSM1.1(m), IPSL.CM5A.LR and MIROC5 of CMIP5 and their equivalents in CMIP6,
331 BCC-CSM2-MR, IPSL-CM6A-LR, MRI.ESM2.0 and MIROC6 have better abilities in
332 replicating historical precipitation properties over Yulin. Projection of precipitation showed an
333 overall decrease in precipitation over Yulin by the GCMs of both CMIPs. The decrease in
334 precipitation would be more in the far future compared to the near future. The expected
335 decreases in precipitation can increase the frequency and intensity of droughts in the area. The

336 findings of the study can be used as a guide in the development of adaptation and mitigation
337 measures against climate change in the area. In the future, more GCMs common to both CMIP5
338 and CMIP6 can be employed when they will be available for CMIP6 for the selection of best
339 performing GCMs. Besides, a more reliable approach can be utilized for the selection of GCMs
340 to avoid the dispute in selection using conventional statistical metrics.

341

342

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346

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478

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489

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491 The authors declare that they have consent to participate in this paper.

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