



Supplementary Figure 1. Primitive Mantle-normalised diagrams for all of the studies samples. A, Dead Sea Fault Zone; B, Gaziantep; C, Karacadağ. A Na-alkaline basalt of Elazığ (Eastern Anatolia, from Di Giuseppe et al., 2017) and a subduction-related basaltic andesite of Kepez Dağ volcanic complex (Central-Eastern Anatolia, from Di Giuseppe et al., 2019) are reported for comparative purposes. Normalisation factors from McDonough and Sun (1995).

Supplementary Table 1. Major Elements, CIPW Norm and Trace elements of collected samples
Supplementary Table 2. Measured and age-corrected data for Sr-Nd-Pb radiogenic isotopes.
Supplementary Table 3. Estimates of Potential Temperature, Pressure and depth of magma segregation

Supplementary Material

The mantle potential temperature and pressure at which magmas segregated from their peridotitic source were estimated using primary melt compositions (e.g., Albarède, 1992; Herzberg et al., 2007; Putirka et al., 2007; Lee et al., 2009). Calculation and results are in [Supplementary Table 3](#). Starting from major element abundances, we calculated the primary magma compositions adding back equilibrium olivine until magmas equilibrates with olivine FO_{90} , using the equations by Pearce (1978), and calculating the composition of equilibrium olivine using K_D $(Fe/Mg)^{ol/liq} = 0.31$ (Putirka, 2005). According to this model, the amount of olivine added to equilibrate starting melt compositions to FO_{90} source was estimated in between 16 and 25% for the selected samples. Then, to calculate potential temperatures and pressures, we used the thermobarometer from Lee et al. (2009). Here, the H_2O content in the melts was estimated by fractionation correction of Ce (ppm), assuming that the magmas emplaced in the study areas have the same H_2O/Ce ratios as oceanic basalts (~ 200 ; Herzberg et al., 2007). The estimated H_2O content was in the range of 0.8 to 2.3 wt.%. Results show that potential melting temperatures (T °C) varies from 1386 to 1528 °C, whereas estimated pressure (P) varies from 1.3 to 3.2 GPa, that is ≈ 38 -97 km of depth ([Supplementary Table 3](#)). Model results are also plotted in Figure 6.

Cited References in Supplementary Materials

- Albarède, F. How Deep Do Common Basaltic Magmas Form and Differentiate? *J. Geophys. Res.* 97, 10.997-11.009 (1992).
- Herzberg, C. *et al.* Temperatures in ambient mantle and plumes: constraints from basalts, picrites and komatiites. *Geochem. Geophys. Geosyst* 8, 1-34 (2007).
- Putirka, K.D., Perfit, M., Ryerson, F.J. & Jackson, M.G. Ambient and excess mantle temperatures, olivine thermometry, and active vs. passive upwelling. *Chem. Geol.* 241, 177–206 (2007)
- Lee, C-T., A., Luffi, P., Plank, T., Dalton, H. & Leeman, W. Constraints on the depths and temperatures of basaltic magma generation on Earth and other terrestrial planets using new thermobarometers for mafic magmas. *Earth Planet. Sci. Lett.* 279, 20-33 (2009)
- Pearce, T.H. Olivine fractionation equations for basaltic and ultrabasic liquids. *Nature* 276, 771-774 (1978).
- Putirka, K. D. Mantle potential temperatures at Hawaii, Iceland, and the mid-ocean ridge system, as inferred from olivine phenocrysts: evidence for thermally driven mantle plume. *Geochem. Geophys. Geosyst* 6, doi:10.1029/2005GC000915 (2005)