

Physical and Chemical Properties of Volcanic Ejecta Produced by the Eruption of Mt. Kirishima Shinmoe-dake, Japan -A Case Study After the Explosive Eruption on March 25, 2018-

Hiromi Akita (✉ akita@bosai.go.jp)

National Research Institute for Earth Science and Disaster Prevention <https://orcid.org/0000-0002-1135-1514>

Research Article

Keywords: Particle size distribution, Infiltration capacity , Soil test , Volcanic ejecta, Volcanic eruption

Posted Date: February 10th, 2021

DOI: <https://doi.org/10.21203/rs.3.rs-169426/v1>

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2 **Kirishima Shinmoe-dake, Japan -A case study after the explosive eruption on**
3 **March 25, 2018-**

4

5 **Hiromi Akita^{a*}**

6 *^a National Research Institute for Earth Science and Disaster Resilience*

7 akita@bosai.go.jp *corresponding author

8

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10 **Keywords:** Particle size distribution; Infiltration capacity; Soil test; Volcanic ejecta;

11 **Volcanic eruption**

12

13 ABSTRACT:

14 The purpose of this study was to clarify the physical and chemical properties of
15 volcanic ejecta that fell after the explosive eruption on March 25, 2018 at Shinmoe-
16 dake. In order to investigate the infiltration capacity, which is one of the physical
17 properties of volcanic ejecta, plots were set up on the outer forest plain, the forest
18 plain, and the forest talus, and a cylindrical frame test was conducted with refer-
19 ence to the method of Takeshita (2011). In addition, soil samples were collected at
20 the three locations where the cylindrical frame test was conducted. The final infil-
21 tration capacity of 38-92 mm / h appeared lower in the forest talus than in the
22 forest plain. It is considered that this is due to the small particle size distribution
23 of 0.1 mm or more, regardless of the particle size of the silt / clay particle size
24 classification. When the chemical properties of the collected volcanic ejecta were

25 examined, Ca and SO₄ were contained at high values. Since these compounds be-
26 come the source of gypsum that reduces infiltration capacity when they react with
27 water, it has become clear that they have the potential to contribute to the gener-
28 ation of debris-flow as a chemical property.

29

30 1. Introduction

31 Mt. Shinmoe-dake in Japan had an explosive eruption for several hundred
32 years from January 26-27, 2011, and volcanic ejecta was deposited in the mountain
33 stream at the foot of the mountain (Oikawa et al. 2013). The deposition of fine-
34 grained volcanic ejecta on slopes is said to significantly reduce the infiltration ca-
35 pacity (Onda et al. 1996), which increases the risk of debris flow generation. At the
36 time of the eruption in January 2011, there was a risk of a debris flow disaster
37 after rainfall, so the Ministry of Land, Infrastructure, Transport and Tourism con-
38 ducted a survey (Ministry of Land, Infrastructure, Transport and Tourism Sabo
39 Department 2011-2018). The eruption at this time entered the lull state in Sep-
40 tember 2011, but an explosive eruption occurred again on October 11, 2017 (Min-
41 istry of Land, Infrastructure, Transport and Tourism Sabo Department 2011-2018).

42 After that, the eruption continued until April 2018, and volcanic ejecta was accu-
43 mulated in the mountain stream at the foot of the mountain. According to Koi et
44 al. (2017), although a sediment outflow was seen, it was directly linked to human
45 damage. No occurrence of debris flow was confirmed. In order to understand the
46 risk of debris flow associated with rainfall after volcanic ejecta fall, it is necessary
47 to conduct a field survey of physical and chemical properties that are soil charac-
48 teristics of volcanic ejecta, unlike the amount of rainfall that can be collected on a
49 desk. However, since the field survey immediately after the eruption is dangerous,
50 the range of people entering is limited, and it is difficult to collect data after the
51 eruption. Therefore, the survey results of volcanic ejecta immediately after the
52 eruption are extremely valuable, and it is expected that accumulating the soil qual-
53 ity data will be academically meaningful and useful for predicting debris flow

54 disasters after the volcanic eruption.

55 Looking at the research examples of the eruption of Mt. Shinmoe-dake in Japan,
56 after the eruption in January 2011, the physical and chemical properties of vol-
57 canic ejecta were investigated at the foot of Mt. Shinmoe-dake, and the relation-
58 ship with the deposition thickness of volcanic ash was considered. (Kisa et al. 2011
59 / Shimizu et al. 2011 / Koi et al. 2013). On the other hand, regarding the eruption
60 after October 2017, although the actual condition of sediment runoff has been in-
61 vestigated by Koi et al. (2017), there are few cases in which the physical and chem-
62 ical properties of volcanic ejecta were investigated. Therefore, the purpose of this
63 study is to clarify the actual physical and chemical properties of volcanic ejecta
64 that fell after the explosive eruption on March 25, 2018 at Shinmoe-dake, Japan.

65

66 2. Methods

67 2.1. Survey site

68 The survey sites are located at the foot of Mt. Shinmoe-dake located in Kiri-
69 shima City, Kagoshima Prefecture (Fig. 1). The ash thickness as of March 6, 2018
70 refers to the result reported by the Ministry of Land, Infrastructure, Transport and
71 Tourism Sabo Department (2011-2018). The research date of this study is March
72 26, the day after the explosive eruption on March 25, 2018. It is assumed that the
73 deposition status of volcanic ejecta will differ depending on the presence or absence
74 of a cover such as a crown on the ground surface. Therefore, it was decided to com-
75 pare inside and outside the forest. For information on the eruption date and scale
76 of Mt. Shinmoe-dake, refer to the Ministry of Land, Infrastructure, Transport and
77 Tourism Sabo Department (2011-2018). For rainfall, I collected the observation

78 values at Yadake Rainfall Observatory of the Ministry of Land, Infrastructure,
79 Transport and Tourism, which is located near the survey site.

80

81 **2.2. Infiltration capacity survey**

82 In order to investigate the infiltration capacity, which is one of the physical
83 properties of volcanic products, three plots are set at each of the three locations
84 (outer forest plain, forest plain, and forest talus), and the method of Takeshita
85 (2011) is referred to. As a result, a cylindrical frame test was conducted (Photo 1
86 (a) - (h)). Cylindrical frame test is a cylindrical frame ($\varphi = 100$ mm, $L = 150$ mm) of
87 about 300 mm inserted into the surface covered with volcanic ejecta, and water
88 injection is continued for a certain period of time. Is to measure. Since this test
89 promotes water infiltration in the direction of gravity by enclosing it in a

90 cylindrical frame, Takeshita (2011) presents the coefficient according to Equation
91 (1) and shows the correction of the real value. In this study as well, the method
92 was quoted and the measurement value of the cylindrical frame was corrected by
93 Equation (1).

$$94 \quad \text{Real infiltration capacity (mm / h)} = 0.24 \cdot \text{cylindrical frame infiltration value}$$
$$95 \quad \dots (1)$$

96

97 It is known that the infiltration into the ground is fast at the beginning of the
98 measurement, but gradually becomes slower in the latter half of the measurement
99 (Takeshita 2011). The test was continued until the permeation rate became slow,
100 and the test was terminated when the change was visually confirmed. In this study,
101 the measured value at this time was defined as the final osmotic permeability

102 value and used for the analysis.

103

104 **2.3. Soil test**

105 In order to investigate the physical and chemical properties of volcanic ejecta
106 at the location where the cylinder frame test was conducted, soil cylinder samples
107 (100 cc) were collected at the three locations plotted at the three locations. For
108 physical properties, soil particle density, ratio of silt and clay, and particle size dis-
109 tribution were measured. The soil test procedure was in accordance with Japanese
110 Industrial Standards (A1202, A1203, A1204).

111 In addition, an X-ray diffraction test was conducted to investigate the type of
112 minerals in the fine-grained volcanic ash deposited on the ground surface. An X-
113 ray diffraction tester (MultiFlex, 40 kV, 40 mA, manufactured by Rigaku Denki

114 Kogyo Co., Ltd.) was used, and measurement was performed using Cu- α rays. Min-
115 eral determination was performed based on the data obtained in this test.

116 The chemical properties of water-soluble ions (Na, K, Ca, Mg, Cl, SO₄, NO₃,
117 NH₄) were measured by the water-soluble amount test, which is one of the chemi-
118 cal analysis methods (JGS 0241-2009). K, Ca, Mg, and Cl were measured with an
119 atomic absorption spectrophotometer (ZA-3000, Hitachi, Ltd.) using the eluate.
120 SO₄, NO₃ and NH₄ were measured using an ion chromatograph (ICA-2000, Toa
121 DKK).

122 For the cation exchange capacity (CEC), add barium chloride to the air-dried
123 sample, centrifuge the sample solution, and read the wavelength of 285.2 nm using
124 an atomic absorption spectrophotometer (Hitachi ZA-3000). It was measured. For
125 exchangeable Ca, use the sample solution (100 mL) filtered by CEC measurement,

126 set the wavelength of atomic absorption spectrophotometer (Hitachi ZA-3000) to
127 422.7 nm, calculate the calcium concentration, and determine the content rate in
128 volcanic ash. The same applies to exchangeable Na, using the sample solution (100
129 mL) filtered by the CEC measurement, setting the wavelength of the atomic ab-
130 sorption spectrophotometer (Hitachi Ltd. ZA-3000) to 589.0 nm, and obtaining the
131 sodium concentration. The content rate was calculated. The ratio of exchangeable
132 ions to cation exchange capacity expressed as a percentage is called ESP (Ex-
133 changeable Sodium Percentage) (Agassi et al. 1981), and can be calculated using
134 Equation (2). The volcanic ejecta are known to show high values, and it is said that
135 in soils of 1.5 or above, chemical crusts are formed with rainfall and the infiltration
136 capacity decreases (Gal et al. 1984).

137
$$\text{ESP}(\%) = (\text{Exchangeable Na}) / (\text{CEC} \times 100) \dots\dots\dots (2)$$

138

139 **3. Results and discussion**

140 **3.1. Eruption status and rainfall status since October 2017**

141 Fig. 2 shows the relationship between the time and scale of eruption from Oc-
142 tober 2017 to April 2018 and hourly rainfall. An explosive eruption occurred on the
143 day before the survey day (March 25, 2018), and it was confirmed that fresh fine-
144 grained volcanic ash had accumulated around the survey site (Photo 1). Since the
145 eruption on October 11, 2017, the 24-hour rainfall has been decreasing until De-
146 cember 2017, and then increasing toward February 2018. Looking at the hourly
147 rainfall, 12 times recorded 10 mm / h or more, and only 3 times recorded 20 mm /
148 h or more. Also, on March 5, 2018, it recorded a maximum of about 36 mm / h
149 during the eruption period. It can be confirmed that there was 10 mm / h or more

150 of rainfall per hour after the eruption, regardless of the time of the eruption, but
151 20 mm / h or more occurred twice (22 mm / h: April 14, 2018, 20 mm / h: April 24
152 of the same year). Therefore, it is estimated that there was no heavy rainfall of 20
153 mm / h or more immediately after the eruption of Mt. Shinmoe-dake.

154

155 **3.2. Physical properties of volcanic products**

156 Fig. 3 shows the time course of the real infiltration capacity. In the forest plain,
157 the infiltration into the ground was too fast and it was not possible to measure it
158 properly, so the comparison is between the forest plain and the forest talus. Both
159 the plain outer the forest and the talus in the forest have been tested three times.
160 Comparing the real infiltration capacity, the talus in the forest tends to be lower
161 than the plain in the forest. In addition, when we look at the time when the real

162 infiltration capacity begins to decline, the outer plains are faster than 1000 sec,
163 while the inner talus tend to be slower. Looking at the 1-hour rainfall data (Fig. 2)
164 at the Yadake Rainfall Observatory of the Ministry of Land, Infrastructure,
165 Transport and Tourism near the survey site, it recorded a maximum of about 36
166 mm / h (March 5, 2018), but on this day Before the explosive eruption. Next, the
167 rainfall with an hourly rainfall is 2 times (22 mm / h: April 14, 2018, 20 mm / h:
168 April 24, the same year), which recorded more than 20mm / h. Comparing the real
169 infiltration capacity at the end of the period after 1000 seconds, it was about 74-
170 184 mm / h on the outer forest plain and about 38-92 mm / h on the forest talus,
171 showing values larger than the hourly rainfall. It became clear that there was.
172 This is consistent with the fact that no debris flow occurred at the foothills of Mt.
173 Shinmoe-dake.

174 Fig. 4 shows the relationship between D_{50} and the final infiltration capacity and
175 soil particle density. The smaller the D_{50} of volcanic ejecta, the higher the soil par-
176 ticle density (Fig. 4 (a)). In addition, the higher the soil particle density, the lower
177 the infiltration capacity tends to be (Fig. 4 (b)). Generally, when the particle size of
178 the soil particles is small, the voids between the soil particles in the soil are small
179 and the density is high, so that it becomes difficult for water to infiltration into the
180 ground. From this, it is considered that the soil particle density increased and the
181 final infiltration capacity decreased because the volcanic ejecta with small grain
182 size were deposited on the forest talus. From this, it is speculated that since the
183 volcanic ejecta with a small grain size are deposited on the talus in the forest, the
184 soil particle density is reduced and the final infiltration capacity is low (Fig. 5 (a)
185 and Fig. 5 (b)). Furthermore, regarding the ratio of silt and clay, when the

186 relationship between D_{50} and soil particle density was examined, no correlation
187 was found between the two. Therefore, regardless of the proportion of fine particles
188 such as silt and clay, it is suggested that the size of soil particles is generally small
189 and that the real infiltration capacity of the end of the talus in the forest has de-
190 creased. Next, Fig. 6 shows the particle size distribution of the collected sample,
191 and Table 1 shows the basic information and measured values. In addition, the
192 grain size of the talus in the forest in this study tends to be smaller than that in
193 the outer forest plain or the forest plain. From these facts, it is suggested that the
194 smaller particle size of the particle size category of 0.1 mm or more than the par-
195 ticle size category of silt / clay may contribute to the decrease of the final infiltra-
196 tion capacity.

197 Fig. 7 shows the relationship between the D_{50} and silt-clay ratios and the final

198 infiltration capacity. Looking at the results of this study, the lower the D_{50} , the
199 lower the final infiltration capacity tends to be ($R = 0.69$, $p = 0.04$). Next, looking
200 at the ratio of silt / clay, it can be seen that the real infiltration capacity at the end
201 of the period tends to decrease slightly as it increases, but the variation is large (R
202 $= 0.50$, $p = 0.17$). This suggests that the D_{50} index, which represents the particle
203 size distribution, may represent a tendency for the final infiltration capacity.

204

205 **3.3. Chemical properties of volcanic products**

206 In order to study the chemical properties, we first clarify the mineral composi-
207 tion of the volcanic products by X-ray diffraction test. Fig. 8 shows the composition
208 of minerals by X-ray diffraction test. First, the mineral that appears in the peak of
209 the X-ray diffraction test of the sample collected in this study is Anorthite. Other

210 minerals detected include Enstatite, Augite, and Ilmenite, but no gypsum that
211 causes a decrease in infiltration capacity when fine-grained volcanic ash solidifies.
212 The survey of this study collected fresh volcanic ash that had fallen after the ex-
213 plosive eruption on March 25, 2018, but there was almost no rainfall within a week
214 before the survey (Fig. 2). It is speculated that no gypsum was detected as a result
215 of no chemical reaction between the deposited volcanic ash and water.

216 Table 2 shows the ion amount and exchangeable base amount of the collected
217 volcanic ejecta. Looking at the amount of detected ions, it can be seen that com-
218 pounds that are the main components of gypsum, such as Ca and SO₄, are con-
219 tained in high values.

220 Looking at the amount of exchangeable base in this study, a high value of 52.5% is
221 shown (Table 2). Onda et al. (1996) calculated ESP by the same method by Mt.

222 Unzen Fugen-dake from June 1990 to June 1995 within the eruption period from
223 November 1990 to March 1995. Mt. Unzen Fugen-dake has frequently had debris
224 flows associated with rainfall since 1991, and this study reported a high value of
225 8.6-76.5% up to a depth of 2 cm on the surface layer of volcanic ash. Comparing the
226 ESP values of this study with the results of Onda et al. (1996), the volcanic ash of
227 Shinmoe-dake may have the potential to reduce infiltration capacity. In connection
228 with this, when water was sprinkled on the forest talus after the cylindrical frame
229 test and the surface of the volcanic ejecta was observed (Photo 2), fine-grained vol-
230 canic ash began to solidify, causing debris flow. It has been confirmed that a surface
231 flow has occurred. From the above, the volcanic ejecta of Shinmoe-dake after the
232 explosive eruption on March 25, 2018 will solidify and become gypsum when there
233 is a large amount of rainfall, and the infiltration capacity will decrease,

234 contributing to the generation of debris flow. It is thought that there is a possibility
235 of doing so.

236

237 **4. Conclusion**

238 This study reveals some of the physical and chemical properties of volcanic
239 ejecta that fell after the March 25, 2018 explosive eruption, albeit with a limited
240 sample size. I would like to continue to accumulate soil data on the physical and
241 chemical properties of volcanic ejecta by conducting similar surveys on volcanic
242 eruptions.

243

244 **Acknowledgment**

245 The hourly rainfall observed at Yadake rainfall observatory near Mt. Shinmoe-

246 dake was provided by Miyazaki river national highway office, Ministry of land,
247 infrastructure, transport and tourism. For the soil test, I cooperated with Urban
248 soil research Co., Ltd. and had a meaningful discussion when interpreting the data.
249 This research analyzed part of the content when the author was in the Public
250 Works Research Institute, and with the help of Mr. Naoki Fujimura of the Public
251 Works Research Institute (at that time), it was greatly supported. Finally, I would
252 like to thank these people.

253

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Figures

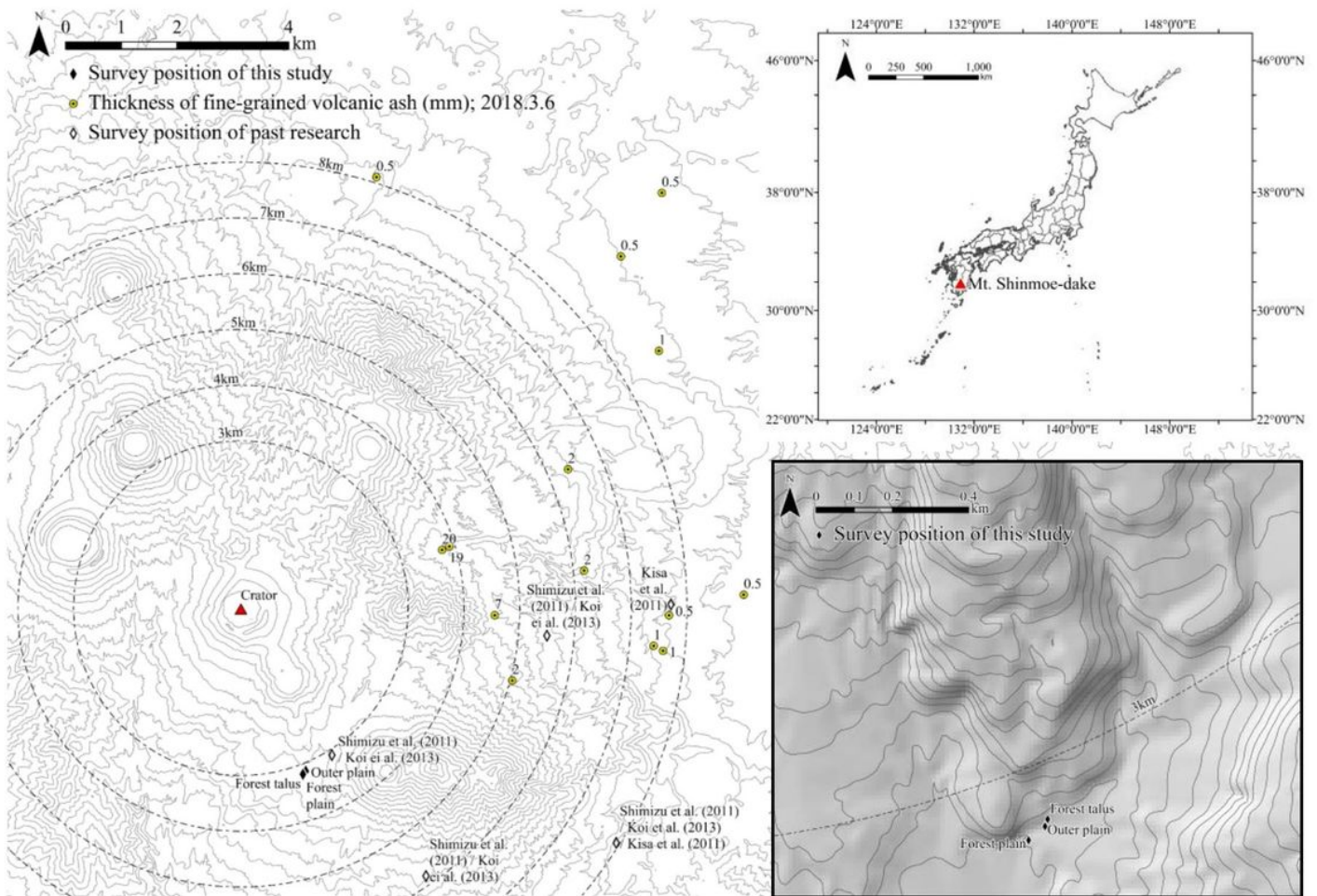


Figure 1

Relationship between distance from Mt. Shinmoe-dake crater and survey location

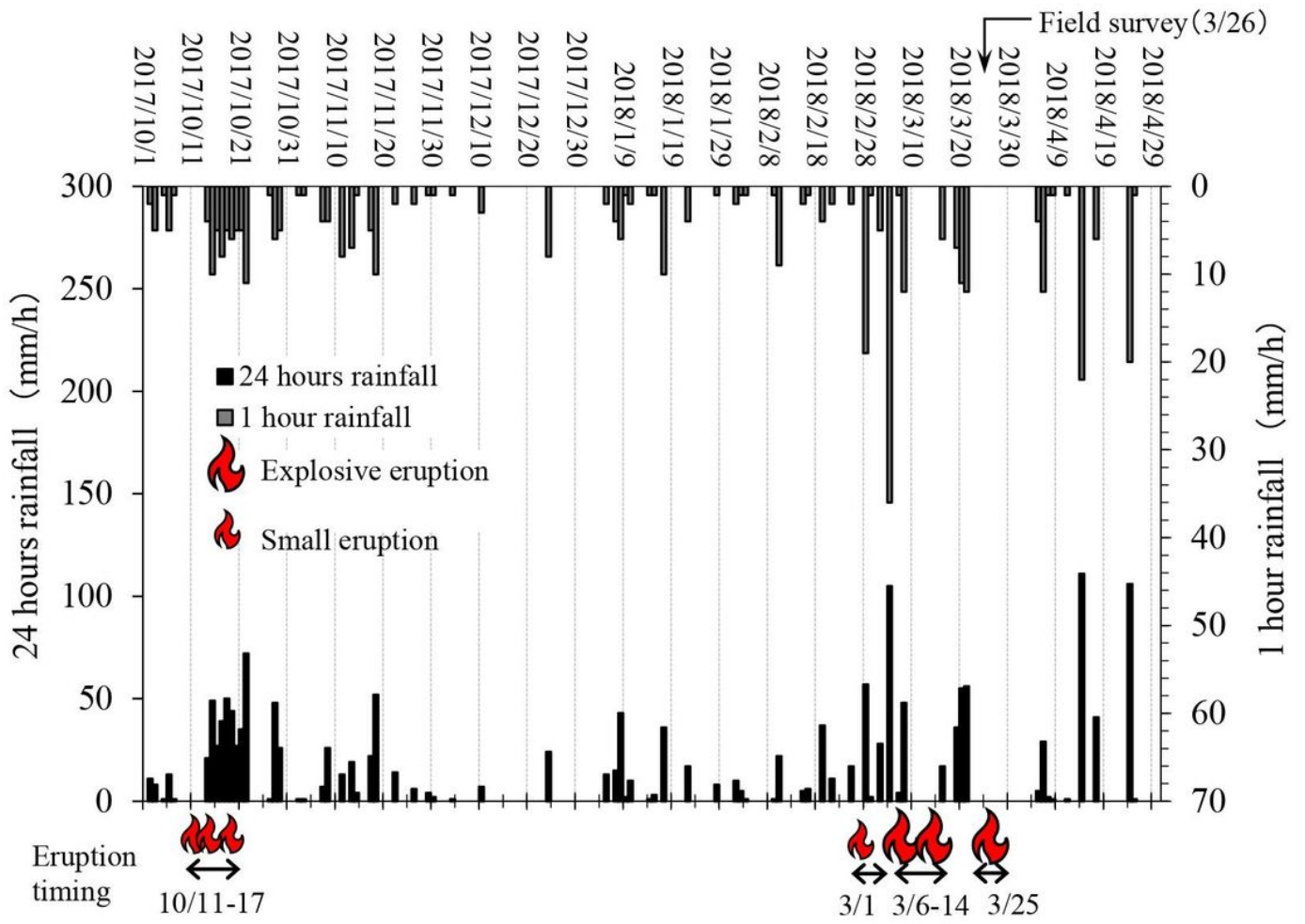


Figure 2

Eruption time of Mt. Shinmoe-dake and hourly rainfall

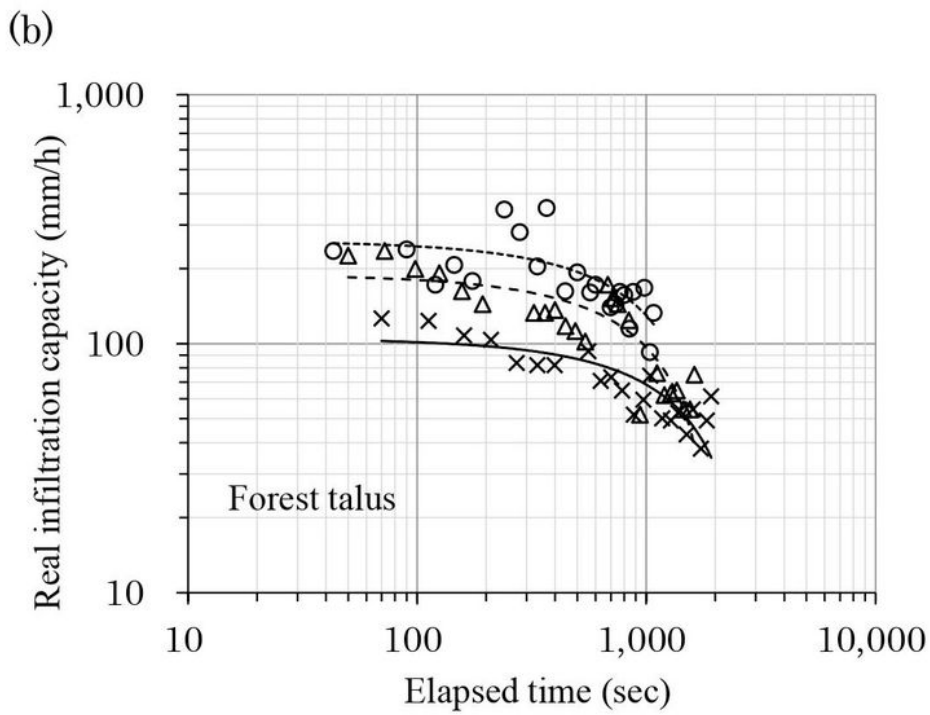
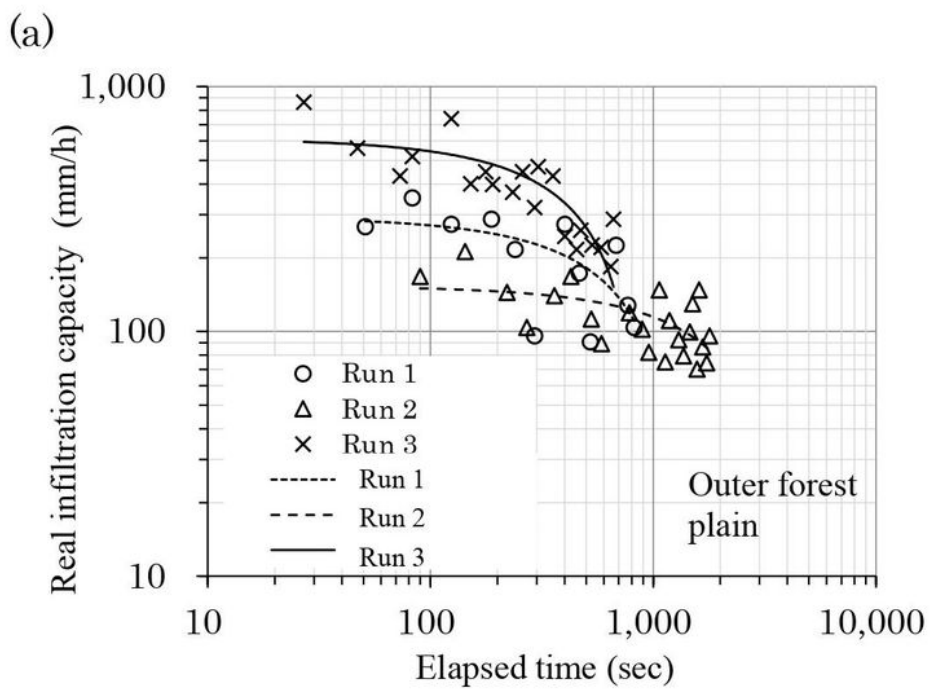


Figure 3

Time-dependent change in real infiltration capacity

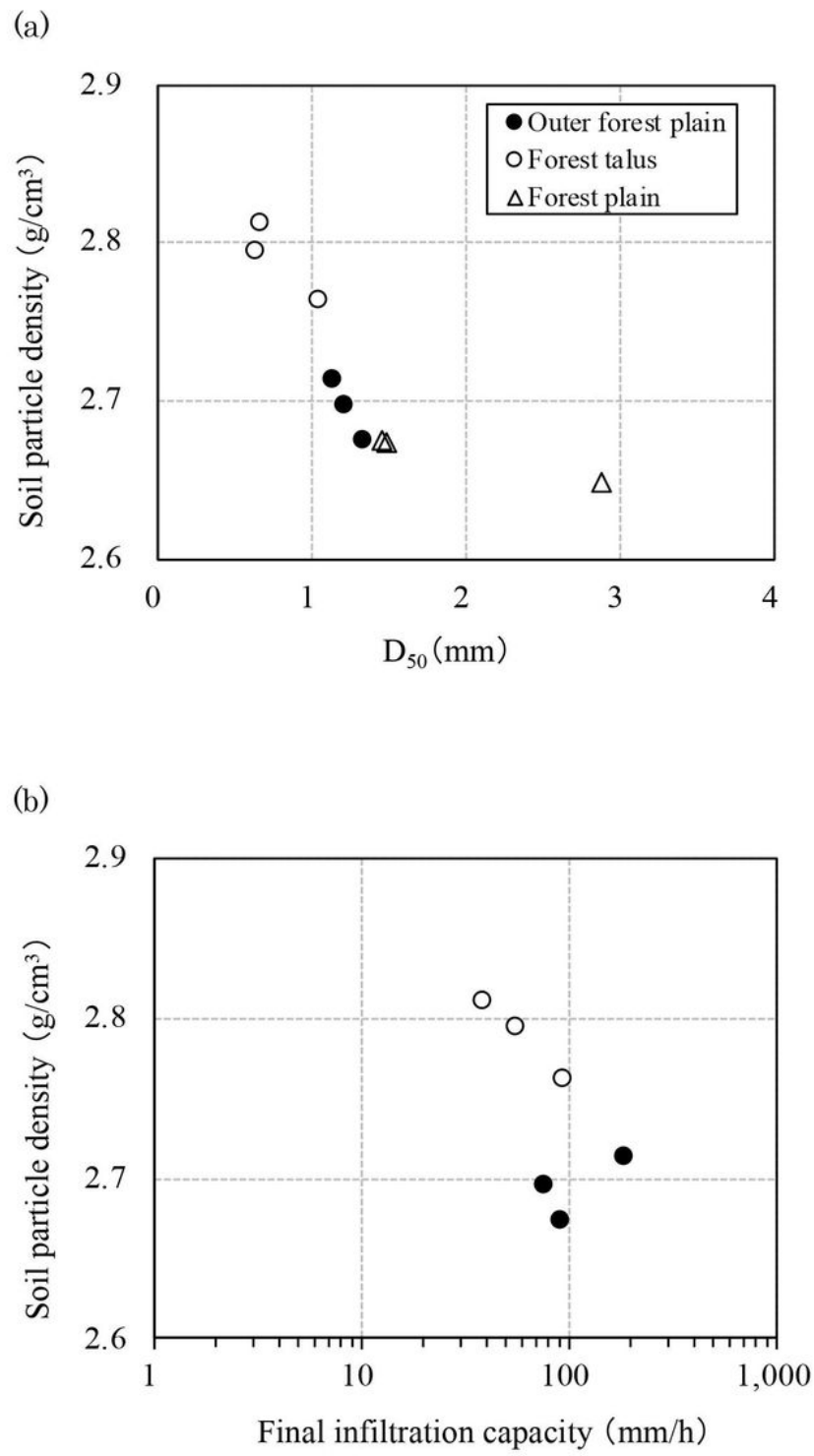


Figure 4

Relationship between D_{50} and final infiltration capacity and soil particle density

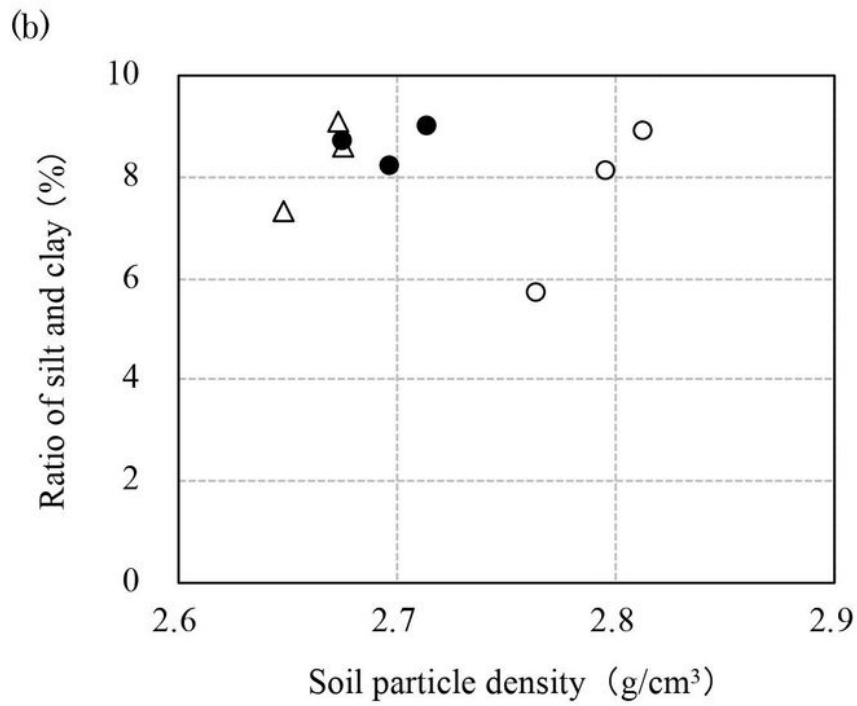
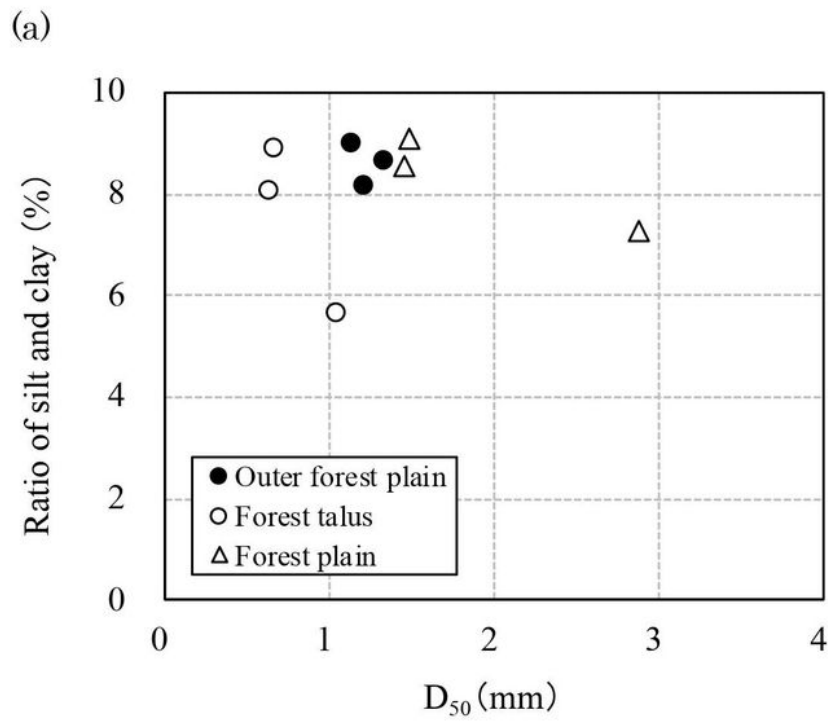


Figure 5

Relationship between D_{50} and soil particle density and ratio of silt and clay

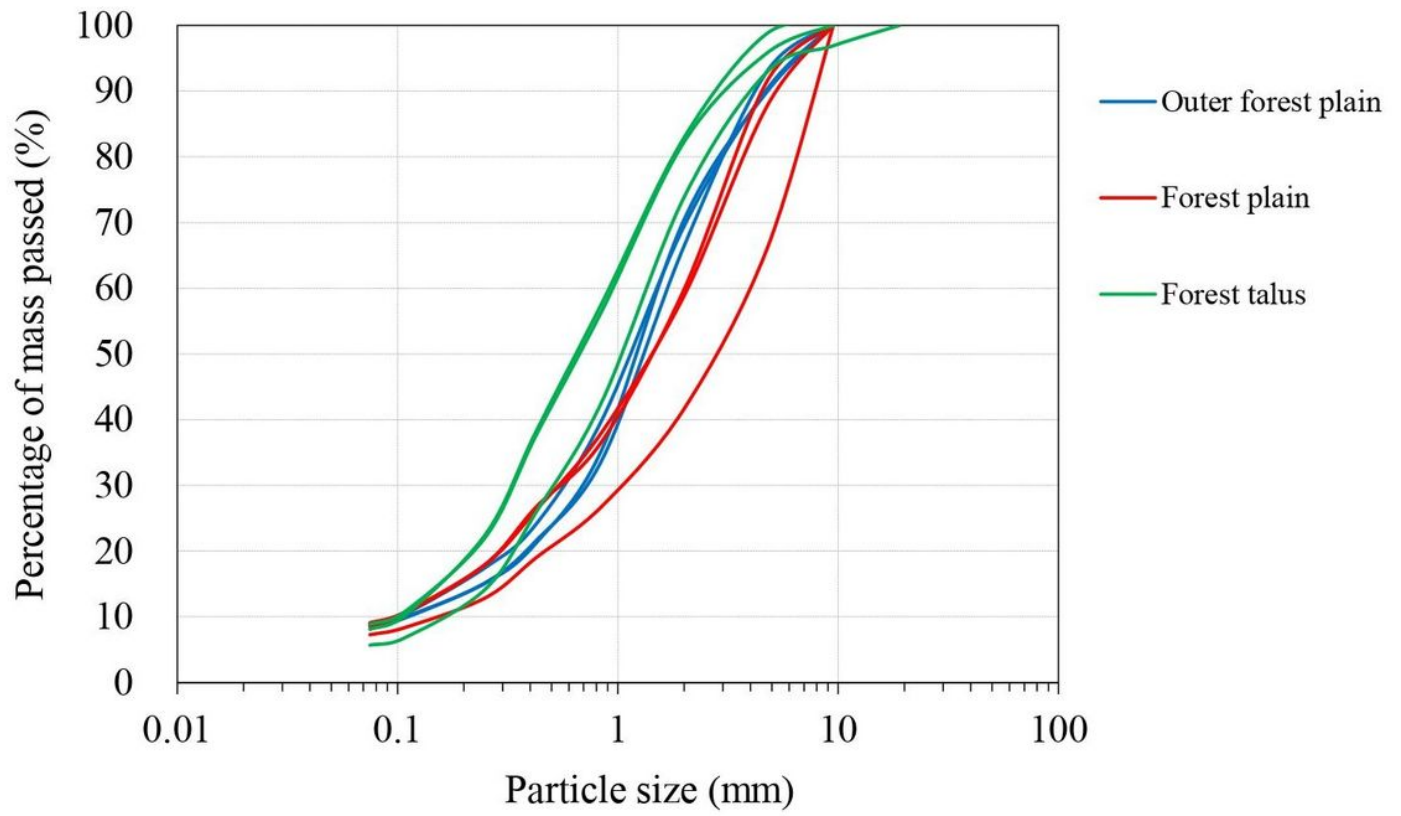
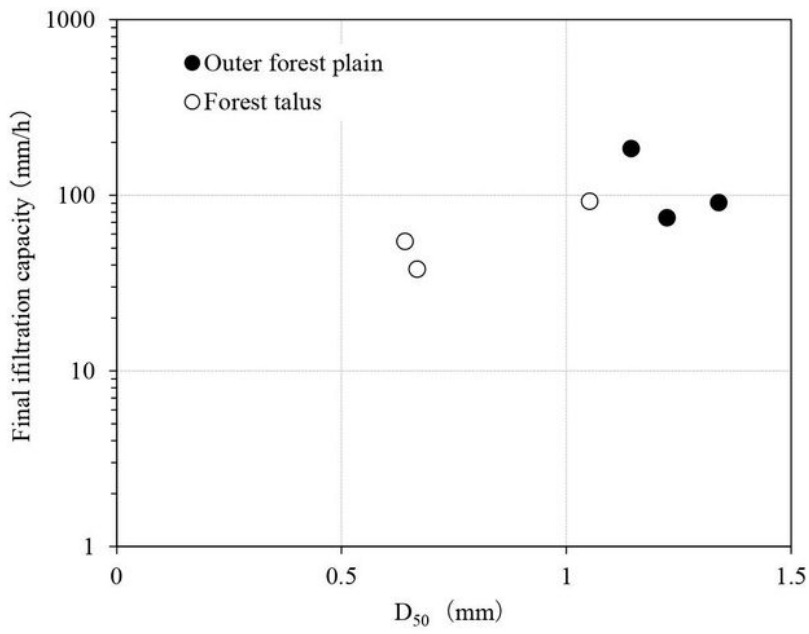


Figure 6

Particle size distribution of collected sample

(a)



(b)

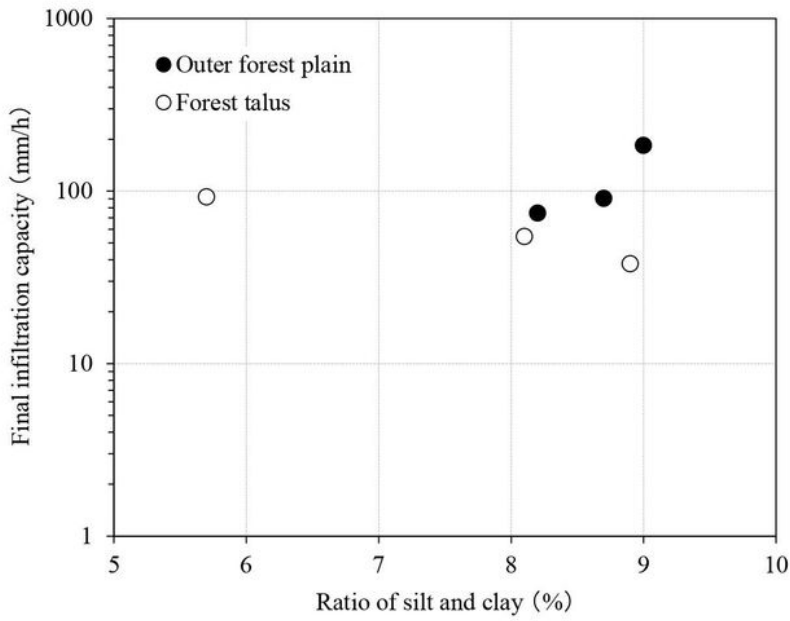


Figure 7

Relationship between D_{50} and the ratio of silt and clay and the final infiltration capacity

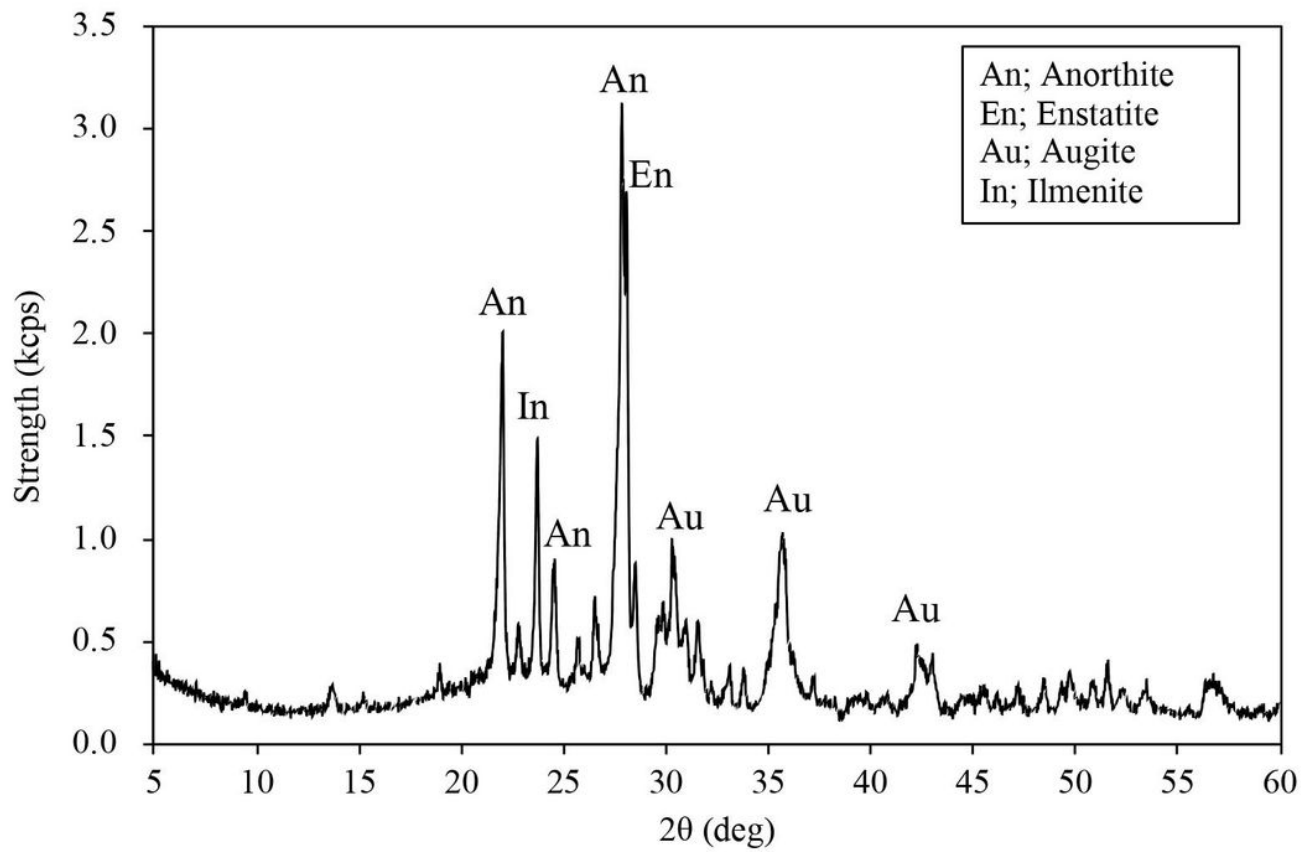


Figure 8

Composition of minerals by X-ray diffraction test

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