

Episodic Magma Hammers in the Recent Cataclysmic Eruption of Hunga Tonga-Hunga Ha'apai

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**Episodic magma hammers for the recent cataclysmic eruption of Hunga
Tonga-Hunga Ha`apai**

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20 **The Hunga Tonga-Hunga Ha'apai volcanic eruption is arguably the most explosive**
21 **since Krakatoa erupted in 1883. Understanding how magma dynamics are regulated**
22 **by volcanic plumbing systems on timescales of seconds to minutes remains**
23 **challenging. Here we identify four similar seismic subevents within a 5-minute**
24 **interval during the intensifying early eruptive phase. Each subevent is similar in**
25 **waveform and duration and is characterized by a sequence of four stages, A-D.**
26 **Initial stage A is marked by an unusual negative P-wave polarity which is best**
27 **explained by an upward, single-force mechanism at the volcano created by a magma**
28 **hammer related to a transiently closed or blocked conduit. Renewed high mass flow**
29 **in the second stage (B) produced a single force down at the volcano which was**
30 **followed by reverberations that represent stages C and D and recovery to the initial**
31 **state. This episodic magma hammer model, which is consistent with thermodynamic**
32 **properties of the multiphase magmatic mixture, yields an estimate of magma mass**
33 **flow in the conduit that is remarkably consistent with discharge into the atmosphere**
34 **estimated from satellite imagery of plume heights.**

35

36 Hunga Tonga-Hunga Ha'apai (HTHH) volcano erupted paroxysmally on 15 January 2022
37 generating a spectacular ensemble of planetary-scale signals¹, including: audible sounds
38 heard in New Zealand and Alaska, the tallest recorded volcanic plume (peak height ~ 58
39 km), >400,000 lighting events, a locally destructive tsunami wave, a global meteotsunami
40 traveling at the speed of sound², and infrasound and seismic waves that circled the Earth

41 multiple times³. In this study, we address the unusual observations of repeating episodic
42 seismic signals during the early stages of the eruption.

43

44 ***Geological Background***

45

46 HTHH is one of many island-arc volcanoes formed as a result of fast (200-250 mm/year)
47 subduction of the Pacific Plate beneath the Indo-Australian plate along the Tonga-
48 Kermadec arc⁴. For HTHH magmas, juvenile H₂O concentrations are ~4-5 wt%, reaching
49 saturation at 5-7 km depths⁵ (see also Supplementary Information Section 2).

50 Crystallization and the concomitant increase in volatile concentrations lead to the
51 development of magma overpressure and the consequent propagation of magma-filled
52 fractures to the surface⁵⁻⁸.

53

54 ***Timings of various volcanic phenomena***

55

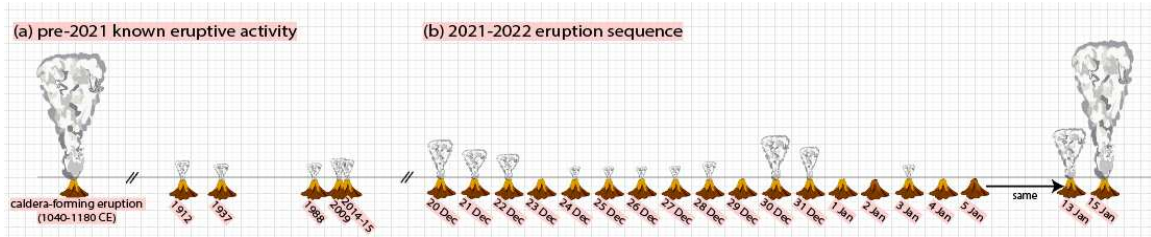
56 Historically, HTHH is a large submarine edifice that has experienced small-scale
57 submarine and Surtseyan activity⁹⁻¹⁴ between large, caldera-forming eruptions⁵ (see
58 Fig.1). The period of activity culminating in the cataclysmic 15 January 2022 eruption
59 began at 1520 UTC on 13 January—producing a 20 cm tsunami and a 20 km tall plume,
60 accompanied by ~190,000 lightning strikes¹⁵. A short-lived land mass between the
61 islands of Hunga Tonga and Hunga Ha’apai formed during the 2014-15 eruption⁹ was
62 partially destroyed by 14 January 2022 activity and totally removed by the 15 January
63 eruption.

64

65 The 15 January volcanic eruption did not onset abruptly. Instead, an ensemble of
66 time-transgressive phenomena overlapped (**Fig. 1c**). The USGS reported a M5.8 seismic
67 event with an origin time of 04:14:45 15 January 2022 (UTC), located at 20.546°S,
68 175.390°E with 0 km depth. This time is not the start time of the eruption, as satellite
69 imagery captured at 0410 UTC shows a rapidly growing volcanic plume already 18 km
70 high—peaking at mesospheric heights of ~58 km at 0430 UTC¹. Extrapolating the
71 infrasound signals in time and distance back to the volcanic location gives a start time
72 around 0402±1 UTC¹, which is more consistent with observed plume heights. Both
73 before and after the M5.8 seismic event, residents of nearby Mango Island reported
74 ashfall persisting throughout the night, suggesting an eruption duration of ~12 hrs.
75 Although the precise timing of the symphony of phenomena remains uncertain, an uptick
76 in lightning intensity around 0412 UTC is consistent with a stronger eruptive pulse
77 starting at 0407-0409 UTC (Fig. 1c), allowing for an estimated five-minute delay
78 between plume ascent and lightning onset, consistent with previous studies of volcanic
79 lightning^{16,17}. The rapid increase in lightning after 0412 UTC is consistent with high mass
80 flow at the vent around 0408 UTC. These phenomena indicate that the eruption was well
81 underway before the 0415 UTC seismic event analyzed in detail below.

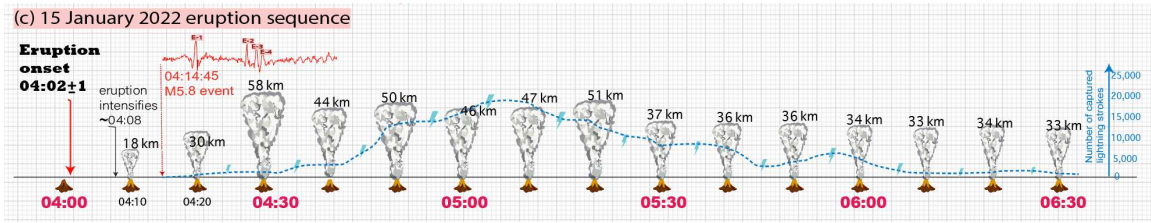
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88 **Figure 1. Timeline of events leading up and including the onset of the 15 January 2022**

89 **eruption of HTHH.** Eruption column heights (labeled by numbers) and eruptive history are

90 reported by the Global Volcanism Program¹⁰⁻¹⁵. For days where no eruptive activity was noted,

91 volcano symbols show no eruption. Days with eruptive activity but no substantial eruption

92 column are marked by lava on the volcano symbols, and eruption columns are to-scale for those

93 days where substantial plume development was reported. Volcanic lightning data are also shown

94 in cyan color¹.

95

96 Although a precise collective timeline remains to be refined, the time of the seismic event

97 studied here is in the crescendoing phase of eruptive activity, correlating with peak vent

98 magma discharge. Based on a mean plume-top upward growth rate of 60 m/s, we

99 estimate the time of peak discharge as ~16 minutes before 0430 UTC (or ~0414 UTC),

100 around the time the seismic event occurred.

101

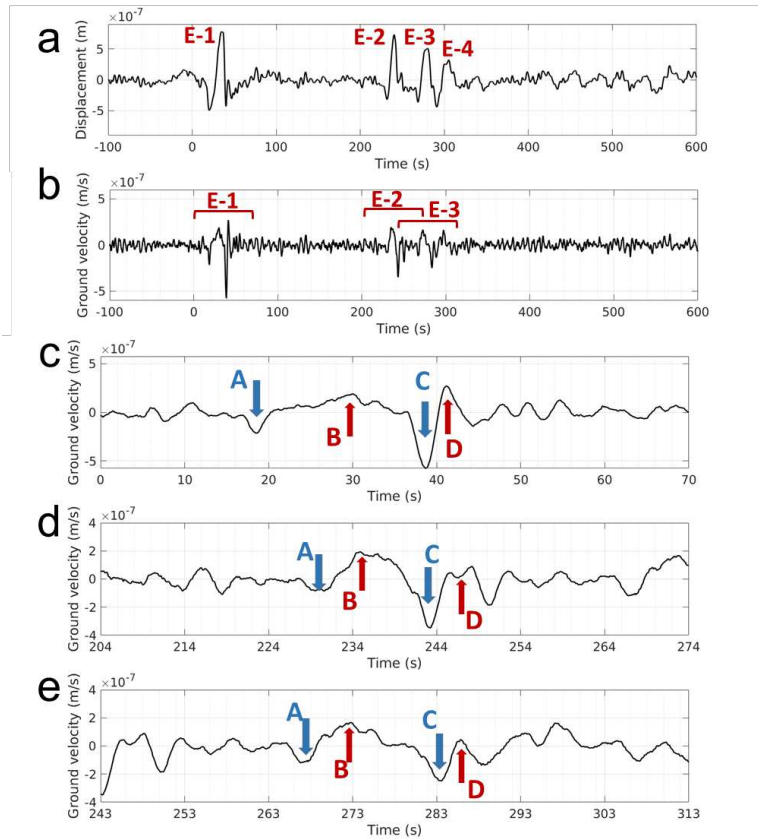
102 *Unusual repeating episodic seismic signals*

103

104 Our seismic analysis involves limited processing of the raw records. We first downloaded
105 teleseismic waveforms recorded at 417 stations worldwide from Incorporated Research
106 Institutions for Seismology (IRIS), removed their instrumental responses, and lowpass
107 filtered below a corner frequency 0.01 Hz. We computed theoretical travel times of the
108 direct *P* waves for all the stations using the USGS reported origin time and location¹⁸.
109 Shifted waveforms were stacked and averaged to obtain the seismic wavelet used to study
110 the source process (**Fig. 2; Supplementary Information Fig. S1**).

111

112 Within a time of ~300 seconds immediately after the reported seismic event, four
113 subevents—E-1, E-2, E-3, and E-4—are visible in the stacked ground displacement
114 seismogram (**Fig. 2a**). Each subevent has a similar duration of ~25 s. If these were
115 regular earthquakes, a 25 s source duration would result in an event of around Mw7.5, far
116 greater than the reported magnitude M5.8 by USGS. The time intervals between the
117 successive subevents are ~204.6 s (E-1&2), ~39.8 s (E-2&3), and ~25.4 s (E-3&4),
118 respectively. We call the subevents E-1 to E-4 ‘episodic’ because their corresponding
119 waveforms are remarkably similar (**Fig. 2c-e**), indicating that similar eruption dynamics
120 likely governed these subevents.



121

122

123 **Figure 2. Source wavelet by stacking global seismic signals.** The vertical-component
 124 seismograms are lowpass filtered below a corner frequency 0.01 Hz; (a) stacked ground
 125 displacement showing four sub-eruptions, E-1 to E-4; (b) stacked ground velocity seismogram;
 126 (c) E-1 velocity waveform; (d) E-2 velocity waveform; (e) E-3 and E-4 velocity waveforms.
 127 Color arrows and numbers in (c)-(e) refer to the eruption stages A-D. Time zero corresponds to
 128 the USGS reported seismic event origin time 04:14:45 Jan 15 2022.

129

130 A puzzling observation is that the polarities of the observed *P*-wave first motions are
 131 found to be downward in both the stacked waveform (**Fig. 2**) and the waveforms
 132 recorded by all global stations at a wide range of azimuths (**Supplementary Information**
 133 **Fig. S1**). Furthermore, similar waveform stacking procedures for global *S* waves using

134 the S wave theoretical travel times yields no visible energy in the transverse component
135 (see **Supplementary Information Fig. S2**). The stacked P and S wave characteristics
136 suggest that the subevents' seismic signals were not caused by earthquake dislocation
137 faulting, but rather by equivalent forces with azimuthal symmetry. Possible forces
138 consistent with the observed first motion are: (1) a single force in the upward (not
139 downward) direction or (2) an implosion/CLVD¹⁹ source at the volcano. Detailed full-
140 wave seismic modeling showed single forces are preferred over either implosion or
141 CLVD (see **Supplementary Information Sec. 1**). Both the implosion and CLVD source
142 near a free surface produce strong P-to-S converted waves not observed in the seismic
143 records (see **Supplementary Information Fig. S3-8**). Our discussions below are based
144 on a single-force model for the seismic source and the modeled force time history (see
145 **Supplementary Information Fig. S9**).

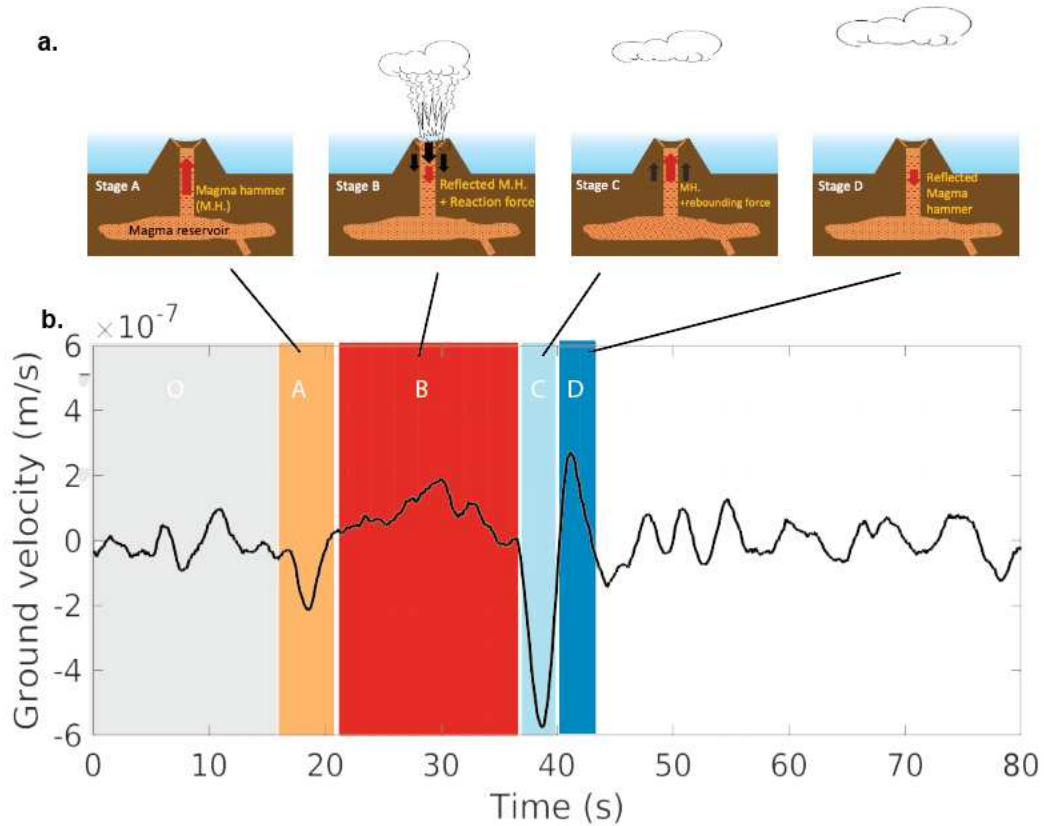
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147 *Eruption stages and magma hammer mechanism*

148

149 Based on the stacked P-wave seismograms (**Fig. 2b-e**), we can decipher four eruptive
150 stages (A, B, C and D) within each of the four subevents. Because each of the seismic
151 subevents exhibits similar waveforms, we focus our analysis on the E-1 time series (**Fig.**
152 **3a**) in the context of the magma transport system, particularly the stage dynamics inferred
153 from the seismic record (**Fig. 3b-g**). The durations for Stages A-D are ~3s, 15s, 3s, and
154 3s, respectively. The timing of E-1A is ~ 4:15:00, 15 seconds after the reported seismic
155 events²⁰.

156



157

158 **Figure 3. Eruption stages and magma plumbing system.** (a) Volcano-reservoir
 159 schematics of states and forces (red arrow: magma hammer; black arrow: reaction force)
 160 corresponding to Stages A-D. Magma reservoir centered at depth of 5-8 km is based on
 161 geobarometry⁵. (b) Globally stacked ground velocity (vertical component) of event E-1.
 162 Time zero corresponds to ~04:14:45 Jan 15, 2022.

163

164 Before Stage-A (time <15s in the record), the stacked seismic record is characterized by
 165 fluctuating ground motion (**Fig. 3b**) likely associated with magma movement in the
 166 conduit, possible phreatomagmatic activity and magma venting. As presented above,
 167 infrasound and lightning data indicate that eruption intensity (the mass discharge at the

168 vent) likely accelerated in the interval 0402-0415 UTC, at which time event E-1
169 occurred.

170

171 ***Stage A: magma hammer and estimates of magma flux and discharge***

172

173 Stage-A (E-1A) lasted about ~ 3.2 s and was characterized by a distinct downward ground
174 motion at the seismometers, meaning that the seismic source was either an upward single-
175 force or a magma reservoir contraction (implosion source) at the volcano. Full-wave
176 seismic modeling (Supplementary Information Section 1) reveals that an upward single-
177 force source, with a force magnitude $\sim 5.0 \times 10^{12} N$ in Stage-A, fits the data far better
178 than an implosion or CLVD source (**Supplementary Information Sec. 1; Fig. S3-S8**).

179 The May 1980 Mount St. Helens (MSH) eruption was also found to produce single forces
180 for seismic wave generation²¹. Unlike the 1980 MSH eruption, ground motion in Stages
181 A and C for all four seismic subevents (E-1~E-4) of the 15 January 2022 HTHH eruption
182 was an upward single force at the volcano. A plausible mechanism to explain this upward
183 force is the water-hammer²²⁻²⁴.

184

- 185 • Magma hammer and piston motion

186

187 In the conventional water-hammer mechanism, water flows steadily through a pipe with
188 an open valve. Rapid closure of this valve creates a surge of water hammer pressure P_H ,
189 acting on the closure which can sometimes burst a steel pipe. Here, we sketch the basic
190 water-hammer concept, modified to the parameters of a multi-phase magma.

191 Mathematically, the water-hammer pressure, P_H , can be described by the well-known
192 Joukowsky equation²² for slow fluid flow ($V \ll c$): $P_H = \rho cV$, where ρ is the fluid
193 density, c is the sound speed in the fluid, and V is the fluid flow speed in the pipe. This
194 equation is also used to compute the pressure exerted on a piston of fluid impacting on a
195 flow stoppage surface. Since magma is involved here rather than water, we apply the
196 term *magma hammer* to describe this process. Because eruptive activity began ~ 10
197 minutes before subevent E-1, a magma transport conduit connected to a surface vent and
198 the plume was likely established. Multiplying both sides of the above equation by the
199 cross-sectional area, A , of the flow conduit, $F_H = P_H A = \rho c Q_V = c Q_m$, where the
200 hammer force, F_H , can be directly estimated by the seismic modeling ($\sim 5 \times 10^{12}$ N for
201 Stage-A), $Q_V = VA$ is the magma volume flux (in m^3/s), and $Q_m = \rho VA$ is the magma
202 mass flux (in kg/s), in the subsurface plumbing system. It is important to note that in this
203 expression the density is that of the magmatic mixture of melt plus fluid. This quantity
204 varies considerably as does the sonic velocity of the magma because the fluid fraction in
205 the mixture increases rather strongly upon decompression as magma flows up the conduit
206 towards the surface vent (see **Supplements Information Sec. 2 and Sec. 3** for details).

207

- 208 • Equation of state and flux estimate

209

210 To apply the water hammer idea to the situation at HTHH, the density and sound wave
211 velocity of the multi-phase magmatic mixture are required. These in turns require
212 consideration of magma composition, degree of volatile saturation, and flow regime.

213 Equilibrium crystallization of last-erupted HTHH andesites was modeled using rhyolite-

214 MELTS²⁵ to approximate the state of an andesitic magma as it undergoes closed-system
215 degassing and ascends adiabatically to the surface from a reservoir depth of ~7.5 km (P =
216 2 kbar). The state of the magma body was recorded at each P - T along an isentropic path
217 and the volume fraction of the exsolved volatiles calculated (**Supplementary**
218 **Information Sec. 2; Fig.S10-11**). The sonic velocity of a two-phase (melt+fluid) mixture
219 was then computed for both bubbly flow (i.e., well mixed) and slug flow (i.e., gas
220 pockets) regimes (see details in **Supplementary Information Sec. 3, Fig.S12-13**). We
221 argue that slug flow is more likely, given the explosiveness of the eruption and short time
222 scales involved. Taking $c \sim 1500$ m/s, we can get a mass flux $Q_m \sim 3.3 \times 10^9$ kg/s using
223 the peak force in E-1A. Approximating the seismic wavelet in E-1A as a triangle, the
224 average force should be halved. Therefore, the average magma mass flux is
225 about $\bar{Q}_m \sim 1.6 \times 10^9$ kg/s. Alternatively, we can estimate the volcanic discharge into
226 the atmosphere²⁶. According to satellite images of the volcanic plume heights, the mass
227 flow to support a 58 km volcanic plume from Yuen et al.¹ is $\sim 1.4 \times 10^9$ kg/s. Considering
228 the uncertainties, the mass flow estimated by our magma-hammer model is of the same
229 order of magnitude as those required to attain peak plume heights.

230

231 Our proposed magma-hammer model is consistent with the seismic data, discharge
232 estimate, physiochemical magma properties, and sonic velocities. We argue that Stage A
233 of E-1~E-4 with similar ground displacements represents magma hammer breaching of a
234 short-lived blockage of the magma conduit, generating an upward single force.

235

236 **Stage B**

237

238 Stage E-1B – a downward single force at the volcano—corresponds mainly to the
239 reactive impulsive force associated with reopening of the magmatic conduit and
240 reestablishment of intense outflow for ~15 s (**Fig. 3a**). The eruption could have sent a
241 series of pulses of volcanic product (ash, pyroclasts, crystals, lithics and seawater-derived
242 steam) into the atmosphere based on observed maximum plume heights referred to
243 earlier. This ‘upside-down rocket’-type eruption created a downward single-force on the
244 order of $\sim 9 \times 10^{12}$ N. The magnitude of the force is comparable to that of the MSH
245 eruption ²¹ of $\sim 10^{13}$ N.

246

247 *Stages C and D*

248

249 Stage E-1C, initiated about 20 s after the onset of E-1A, exhibits an upward force of
250 $\sim 4 \times 10^{12}$ N, which we associate with a second-phase closing of the conduit and
251 subsequent generation of a magma-hammer upward force (details in **Supplementary**
252 **Information Sec. 3**). Stage E-1C (**Fig. 3a**) and Stage E-1D (**Fig. 3a**) do not mirror or
253 repeat the Stage-A to Stage-B pattern because E-1D is only ~3s but E-1B is ~15s. Stage
254 E-1B exerted a sustained downward force, probably depressing both the magma in the
255 conduit and the magma chamber and surrounding rock. It is likely that Stage E-1C is a
256 magma-hammer force related to the backflow impacting the overburden rock. E-1D
257 represents a reflected magma-hammer signal traveling down to impact the magma
258 reservoir that created a downward force. Stages E-1C and E-1D represent one cycle of
259 oscillatory piston motion as magma travels through the conduit, from the reservoir to the

260 surface. The ~6s period corresponds to the two-way travel time of the magma-hammer
261 signal in the conduit. Taking an estimated average traveling sonic velocity of ~2,000 m/s,
262 we can determine the depth of the magma chamber at ~6 km, consistent with
263 geobarometric estimates ⁵.

264

265 **Repeating episodic seismic signals**

266

267 Stages A-D for the E-1 event were repeated three more times (Fig.2), each of which can
268 be interpreted similarly to the E-1A to E-1D sequence of UP-DOWN-UP-DOWN forces,
269 by virtue of similar waveforms. These large magnitude forces are short in duration and
270 could produce loud audible sounds as recorded by observers and infrasonic sensors
271 around the world¹. The timings of the seismic energy envelopes of E-1~E-4 and those of
272 the air pressure waveform recorded by Station MSVF in Fiji are in good agreement ¹.

273

274 ***Conclusions***

275

276 The HTHH eruption will be an important event in volcanology and other geosciences. It
277 is unusual that a volcanic eruption could generate such energetic P waves with negative
278 polarities worldwide. It is even more astonishing that HTHH has a sequence of four
279 similar episodic events, separated in time by only tens of seconds, in the early
280 intensifying phase of the eruption. It reveals a rapid but regulated nature of the shallow
281 magma transport system manifested as repeating magma hammers due to brief (order
282 seconds) cessation of magma flow in the conduit. Our scenario is consistent with

283 available seismic, infrasound, lightning, and satellite data. It is also consistent with the
284 mechanics and thermodynamics of magma under subsurface conditions accounting for
285 the variation in the volume fraction of fluid—mainly supercritical H₂O—in the magma.
286 Examination and modeling of the seismic data suggest that HTHH produced the largest
287 known magma-hammer force associated with a violent volcanic eruption.

288

289 The eruptive volume of the epic 15-Jan-2022 HTHH event is presumably a small
290 portion of the total available magma in the reservoir that, in turn, is repeatedly recharged
291 by additions from below. Future work will benefit from new 3D marine seismic imaging
292 of the plumbing system, with the goal of imaging not only the magma reservoir at 5-8 km
293 depth but also the size and geometry of the plexus of magma conduits to further test our
294 model and perhaps constrain the timing, duration, and magnitude of a future catastrophic
295 eruption using the episodic nature of the system.

296

297 **STOP STOP STOP STOP STOP STOP STOP STOP**

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394

395 **Data availability**

396 Seismic data are downloaded from IRIS. The facilities of IRIS Data Services, and
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402

403 **Author Contribution Statement**

404 From ~February 1 to May 3, the authors held weekly Zoom meeting discussions. DAY is
405 instrumental in forming the team. This manuscript is a result of this effort. An initial
406 manuscript draft was provided by YZ. All the authors contributed to writing. YZ
407 identified the 4 episodic events. HH did seismic modeling (SI Sec. 1) and Fig. 2. FJS led
408 the effort in the magma hammer modeling in SI Sec. 3 and plume height analysis. MS led
409 geological/petrological evaluation of HTHH, the timeline in Fig. 1, and rhyolite-MELT's
410 modeling of magma mixture (in SI Sec. 2). GT led the infrasound and lightning record
411 analysis for the eruption timeline. TJL and YJ suggested the water hammer idea. SRM
412 contributed to eruption sequence/timeline analysis. KM studied tsunami records and
413 contributed to the eruption timeline. ZGP contributed to the seismic analyses. All authors

414 participated in the weekly discussions which have shaped the manuscript to its current
415 form.

416

417 **Competing interests**

418 The authors declare no competing interests.

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